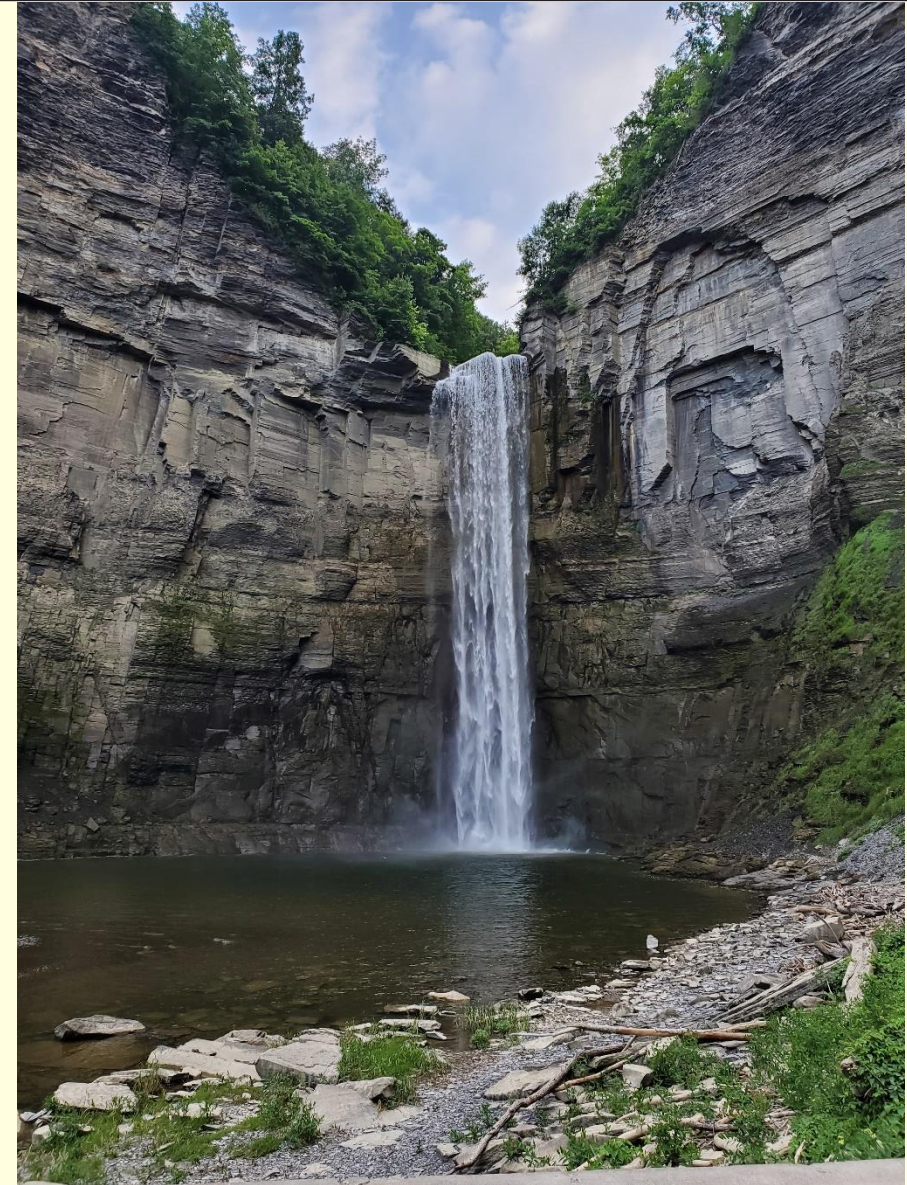


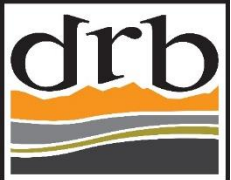
Inorganic Elemental Data 101: Evaluating Redox Conditions in Mudstones

Randy Blood



- **Not all black shales accumulate under the same depositional conditions (applies to the basin and stratigraphic scale for a particular mudstone as well)**
 - Deep stratified basins, shallow energetic environments, coastal upwelling, restricted lagoon, lacustrine
- **Not all elemental interpretations are straight forward and they should be used in the context of the entire data set**
- **Likewise elemental data sets should be combined with other data sets to understand and predict reservoir facies distribution**
 - Standard core analysis, sedimentological analysis, field studies

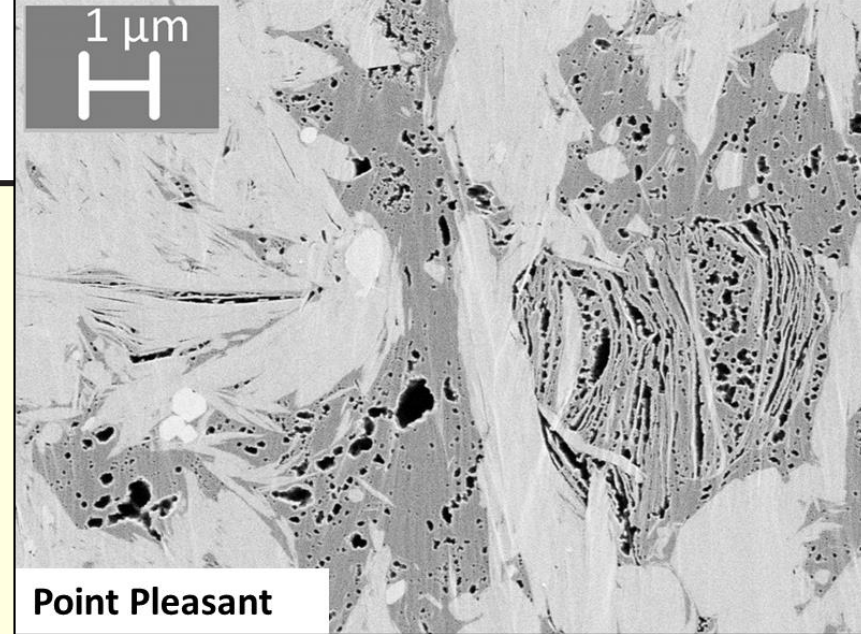




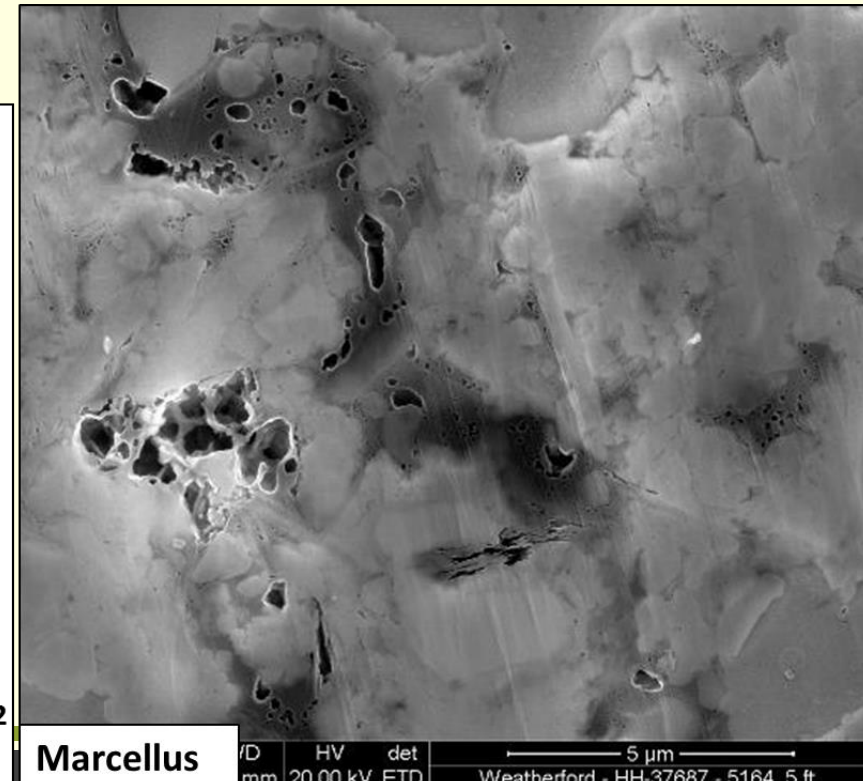
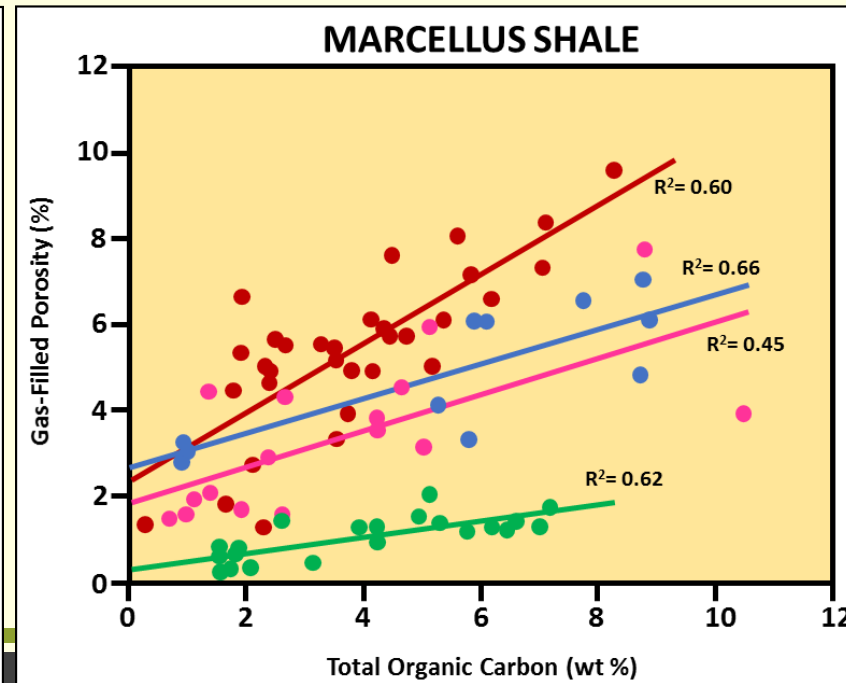
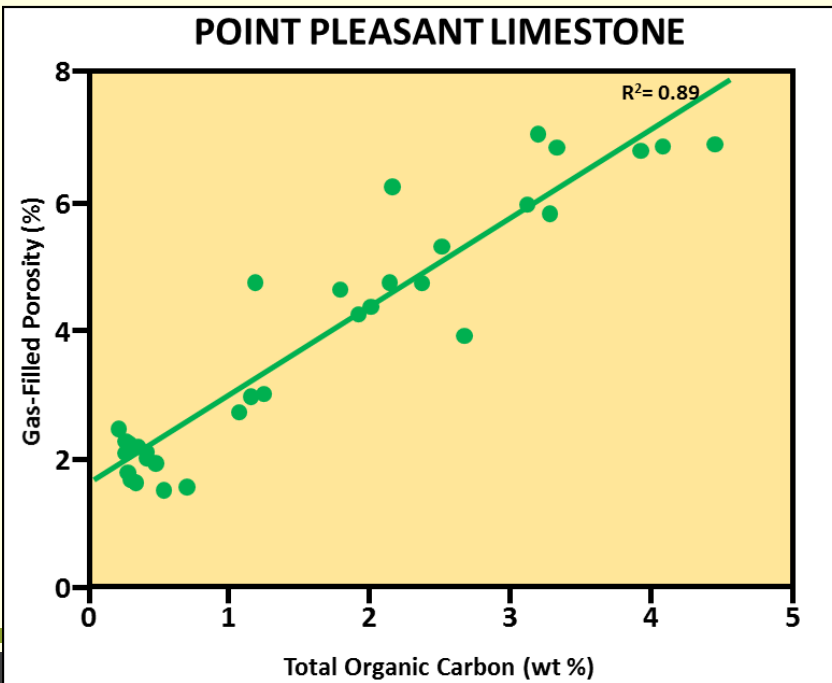
Purpose of the Work

Production Mechanisms

- In many unconventional reservoirs, organic matter hosts the majority of porosity.
- Indeed, some reservoirs depict a strong correlation between gas-filled porosity and TOC.
- Reservoirs are self sourcing, the original TOC represents the starting material for hydrocarbon generation.



Point Pleasant



Marcellus

WD HV det
mm 20.00 kV ETD
5 μm
Weatherford - HH-37687 - 5164. 5 ft.



Purpose of the Work

Visual Characterization of Organic-rich Mudstone

UTICA

**POINT
PLEASANT**

MARCELLUS

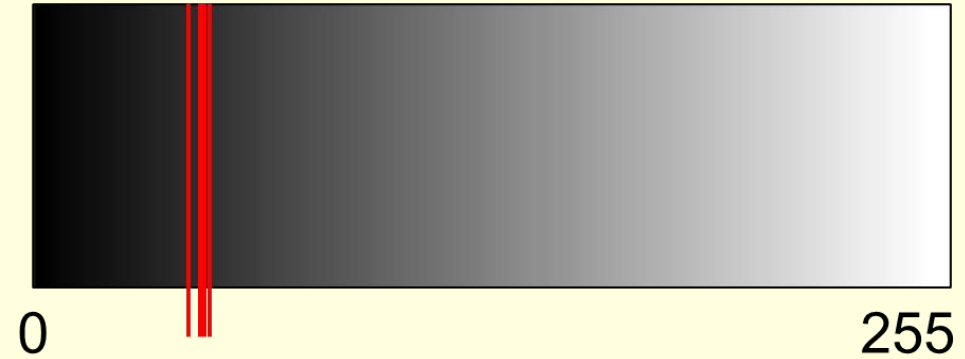
GENESEO

Average color:
46.9

Average color:
49.8

Average color:
47.2

Average color:
42.4

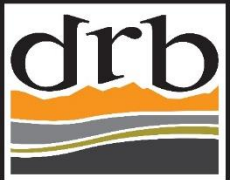


Quantitatively there is a <3% difference in the color of these formations.

These rocks range from oxic/dysoxic to anoxic/euxinic.

Visual inspection alone is not enough to understand the depositional environment.

While extremely useful, even sedimentary structures (bedding laminae, erosion surfaces, fossils, etc.) cannot definitively inform on the redox state of rocks.

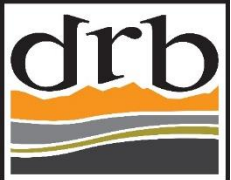


Water oxygenation conditions

Oxygen (ml/l)	Oxygenation regime environments	Biofacies
>2.0	Oxic	Aerobic
2.0-0.2	Dysoxic	Dysaerobic
2.0-1.0	moderate	
1.0-0.5	severe	
0.5-0.2	extreme	
0.2-0.0	Suboxic	Quasi-anaerobic
0.0	Anoxic	Anaerobic
0.0 (H ₂ S)	Euxinic	

Modified from Tyson and Pearson, 1991

- Three chemical species play a role in the oxidization of organic matter:
 - Oxygen
 - Nitrate
 - Sulfate
- Eukaryote and prokaryote communities use oxygen to oxidize organic material until the supply of oxygen is dissolved.
- Denitrifying bacteria then consume nitrate until it is dissolved.
- The reservoir of sulfate is used to oxidize organic matter by sulfate-reducing bacteria.
- Finally, organic matter can undergo fermentation before preservation.
- These processes further influence the degree to which certain elements remain in the water or accumulate in sediment.
- These are known as Redox Sensitive Trace Elements (RSTE).



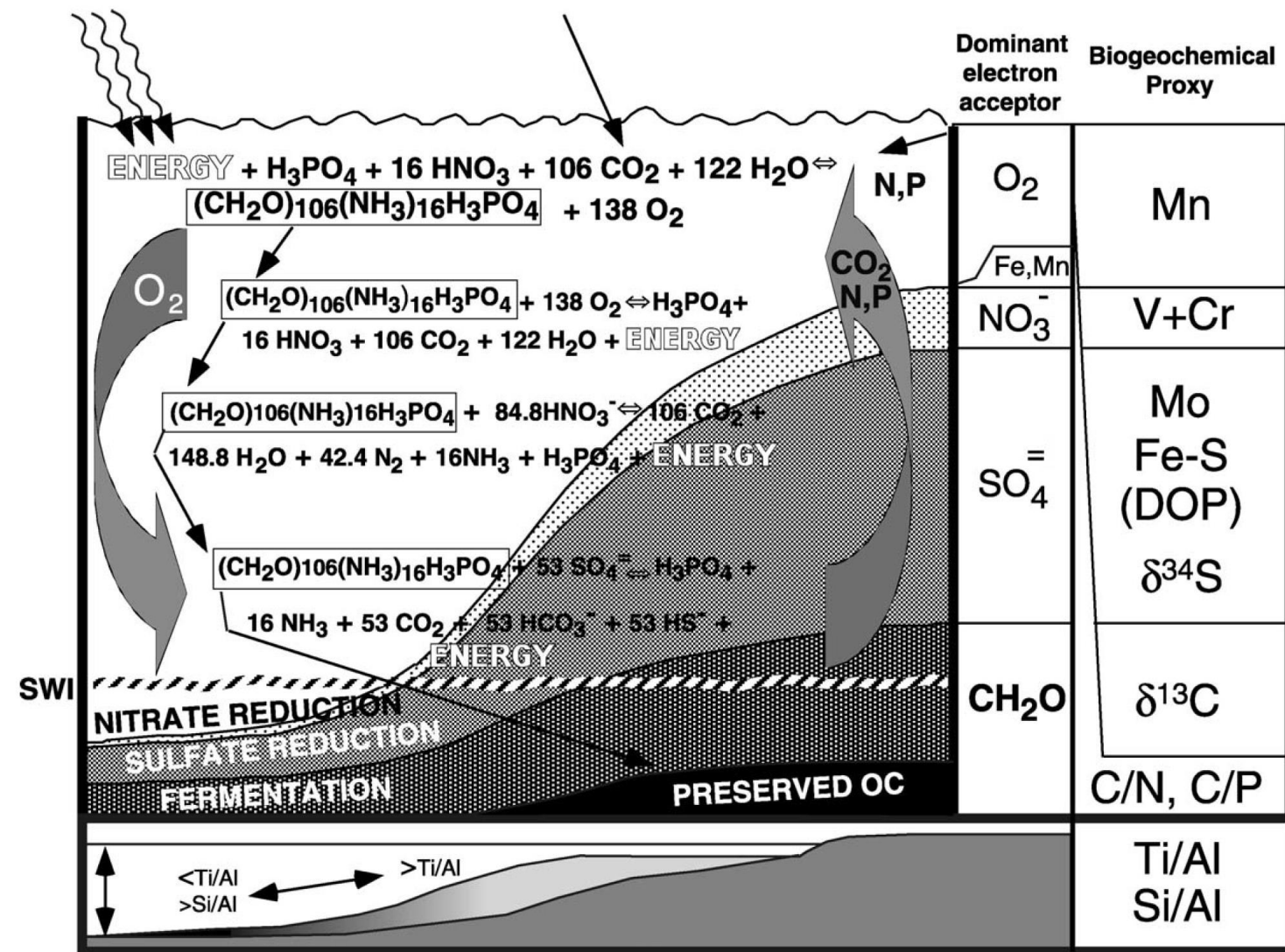
Redox sensitive trace elements

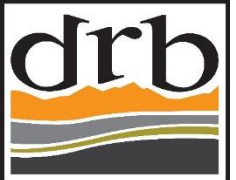
1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	
hydrogen 1 H 1.0079																	helium 2 He 4.0026	
lithium 3 Li 6.941	beryllium 4 Be 9.0122											boron 5 B 10.811	carbon 6 C 12.011	nitrogen 7 N 14.007	oxygen 8 O 15.999	fluorine 9 F 18.998	neon 10 Ne 20.180	
sodium 11 Na 22.990	magnesium 12 Mg 24.305											aluminium 13 Al 26.982	silicon 14 Si 28.086	phosphorus 15 P 30.974	sulfur 16 S 32.065	chlorine 17 Cl 35.453	argon 18 Ar 39.948	
potassium 19 K 39.098	calcium 20 Ca 40.078											gallium 31 Ga 69.723	germanium 32 Ge 72.61	arsenic 33 As 74.922	selenium 34 Se 78.96	bromine 35 Br 79.904	krypton 36 Kr 83.80	
rubidium 37 Rb 85.468	strontium 38 Sr 87.62											indium 49 In 114.82	tin 50 Sn 118.71	antimony 51 Sb 121.76	tellurium 52 Te 127.60	iodine 53 I 126.90	xenon 54 Xe 131.29	
caesium 55 Cs 132.91	barium 56 Ba 137.33	57-70 *										thallium 81 Tl 204.38	lead 82 Pb 207.2	bismuth 83 Bi 208.98	polonium 84 Po [209]	astatine 85 At [210]	radon 86 Rn [222]	
francium 87 Fr [223]	radium 88 Ra [226]	89-102 **										ununbium 112 Uub [285]	ununtrium 113 Uut [284]	ununquadium 114 Uuq [289]	ununpentium 115 Uup [288]	ununhexium 116 Uuh [293]	ununseptium 117 Uus —	ununoctium 118 Uuo [294]

*lanthanoids

**actinoids

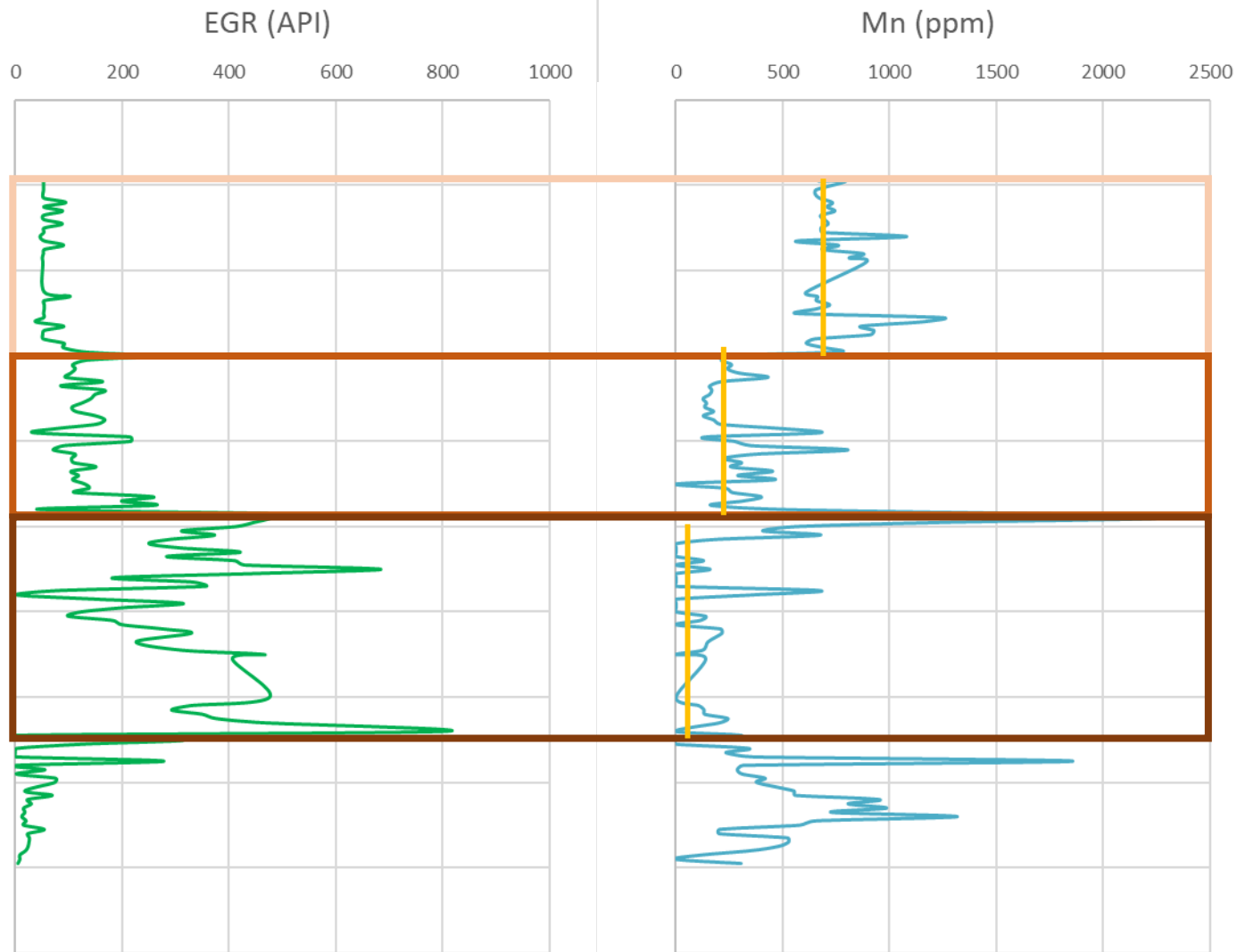
lanthanum 57 La 138.91	cerium 58 Ce 140.12	praseodymium 59 Pr 140.91	neodymium 60 Nd 144.24	promethium 61 Pm [145]	samarium 62 Sm 150.36	europium 63 Eu 151.96	gadolinium 64 Gd 157.25	terbium 65 Tb 158.93	dysprosium 66 Dy 162.50	holmium 67 Ho 164.93	erbium 68 Er 167.26	thulium 69 Tm 168.93	ytterbium 70 Yb 173.06
actinium 89 Ac [227]	thorium 90 Th 232.04	protactinium 91 Pa 231.04	uranium 92 U 238.03	neptunium 93 Np [237]	plutonium 94 Pu [244]	americium 95 Am [243]	curium 96 Cm [247]	berkelium 97 Bk [247]	californium 98 Cf [251]	einsteinium 99 Es [252]	fermium 100 Fm [257]	mendelevium 101 Md [258]	nobelium 102 No [259]





Geochemical proxies for redox conditions

Manganese (Mn) and oxygen



Ludlowville Shale

Dominantly oxic gray shale

Levanna Shale

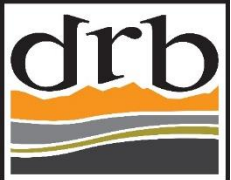
Dysoxic-anoxic black shale

Marcellus Shale

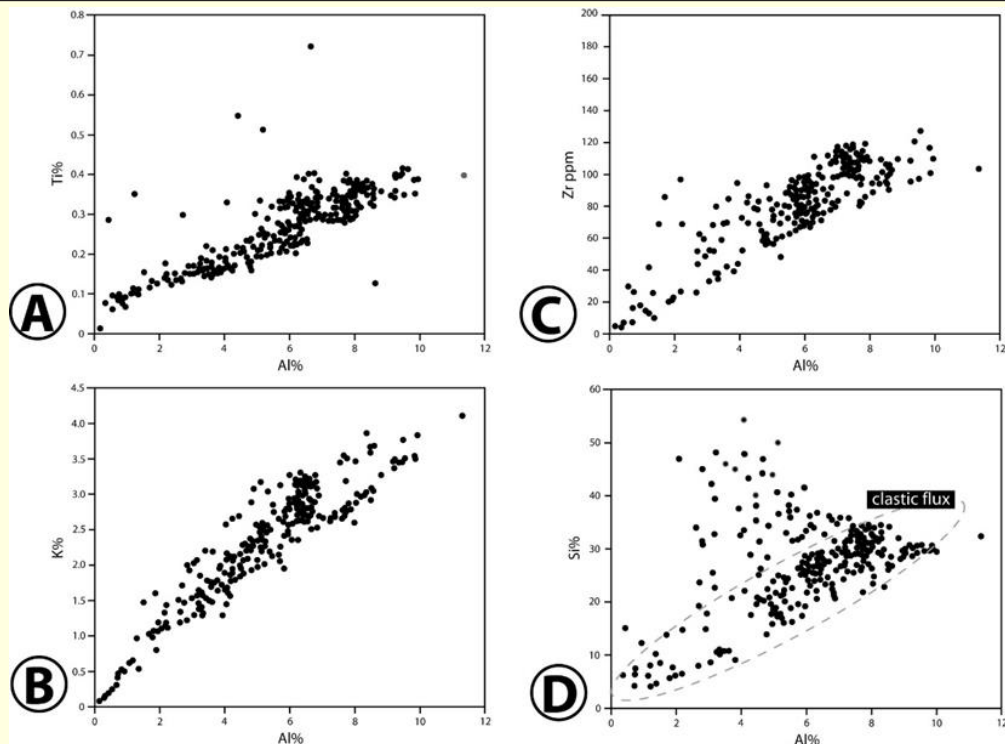
Anoxic-Euxinic black shale

Onondaga Limestone

Oxic limestone



Al normalization and enrichment factors



- Al cross plotted against clastic-derived elements Ti, Zr, and K.
- Note that while much of Si defines a clastic trend, the relationship is more nuanced.
- While Al, Ti, Zr, and K demonstrate positive covariance, their relationships to Al (dominantly a signal for clay) provides insight into the grainsize and energy of sediment delivered to the basin.

$$\text{Enrichment Factor (EF)} = \frac{\text{Element}_{\text{sample}} / \text{Al}_{\text{sample}}}{\text{Element}_{\text{average shale}} / \text{Al}_{\text{average shale}}}$$

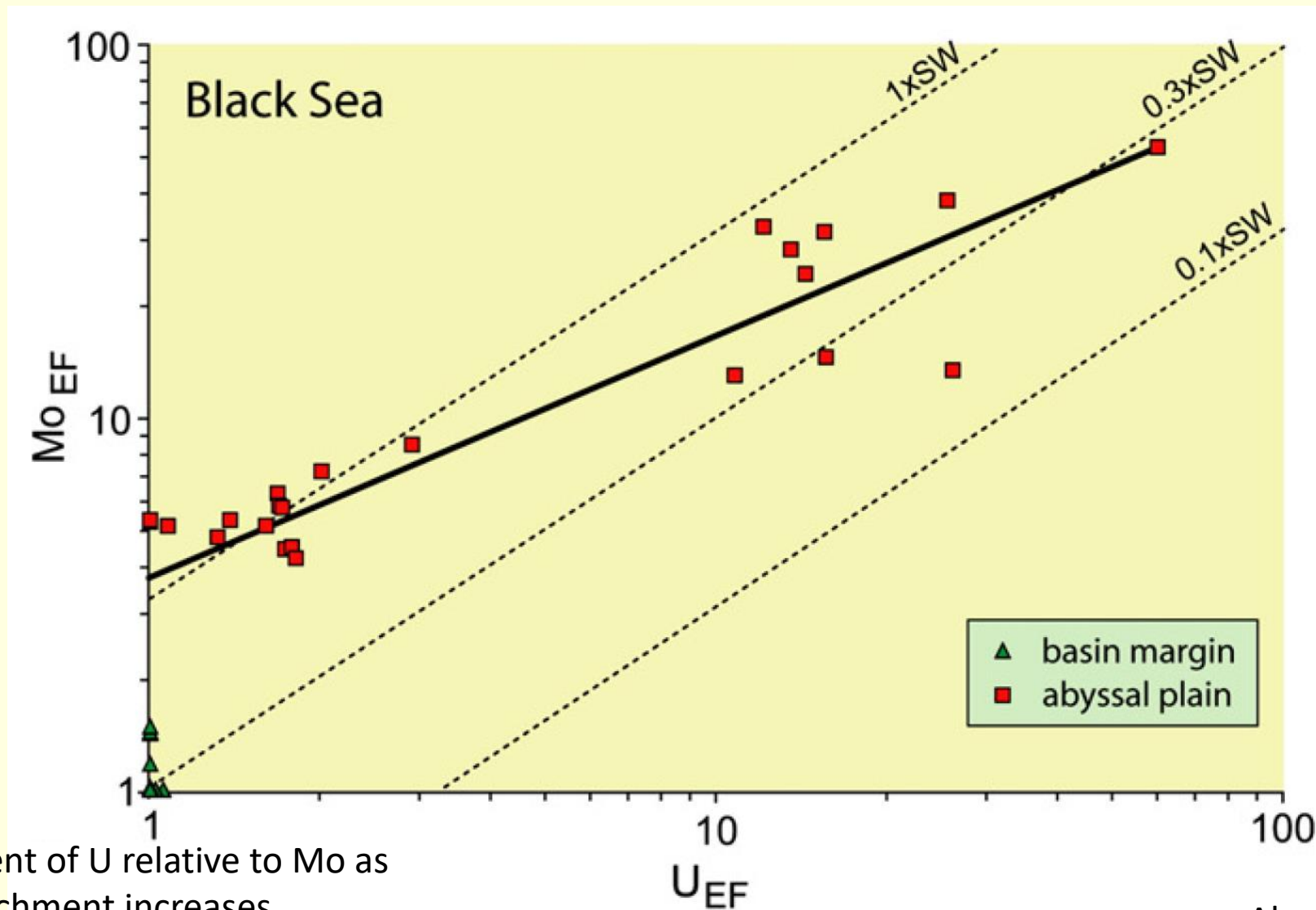
- Elemental data commonly cast as an enrichment factor.
- The element is normalized to the Al content of the sample
 - Accounts for an increase in abundance due to increased sediment supply
- The element/Al is then normalized to the average shale value (Wedephol 1971, 1991).
 - Unity implies elemental abundance is typical of shales
 - $\text{EF} > 3$ implies significant enrichment of that element
 - $\text{EF} < 1$ suggests depletion of that element

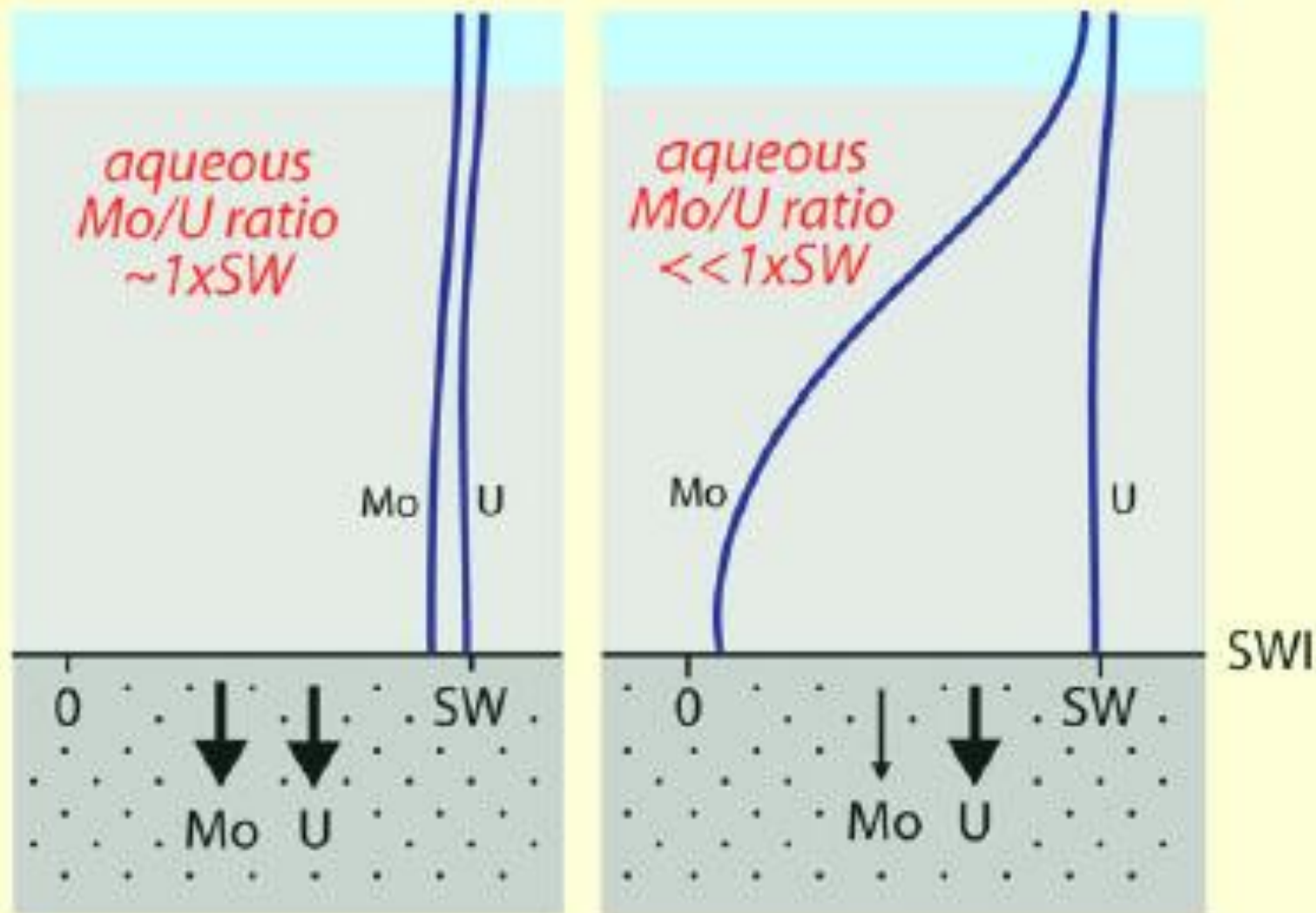
- Mo and U
 - Redox sensitive metals especially useful for paleoenvironmental and hydrographic studies
- Present in low concentrations in the upper crust-
 - Mo ~3.7 ppm
 - U ~ 2.7 ppm (Taylor and McLennan, 1985)
- Both exhibit conservative behavior under oxic conditions;
- Both elements have long residence times in seawater
 - Mo ~0.78 MY
 - U ~0.45 MY
- Both exhibit roughly equal concentrations in seawater globally:
 - Mo/U = 7.53 molar ratio – Pacific Ocean
 - Mo/U = 7.90 molar ratio – Atlantic Ocean
- Both exhibit low concentrations in plankton – enrichment in sediment can be related to authigenic uptake from seawater;



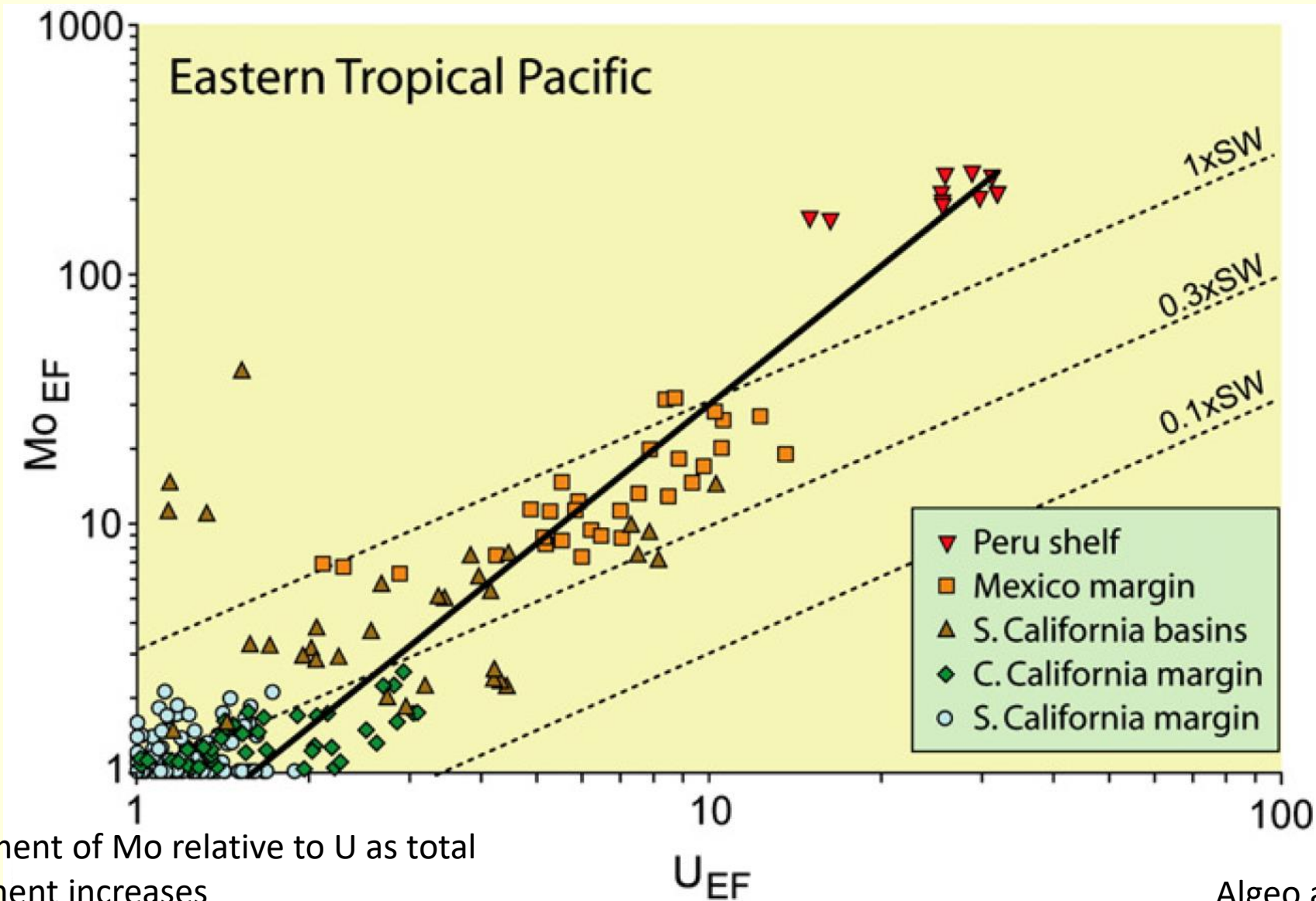
- Significant differences
- U uptake is linked to organic carbon content (TOC) and occurs by diffusion across the sediment-water interface...organic matter serves as a sorbant...sedimentation rate is very important to this
- Mo uptake by sediment requires the presence of H_2S (euxinic conditions) and can be accelerated by particulate Mn-Fe-oxyhydroxide shuttle (U is unaffected by this); less covariance with TOC



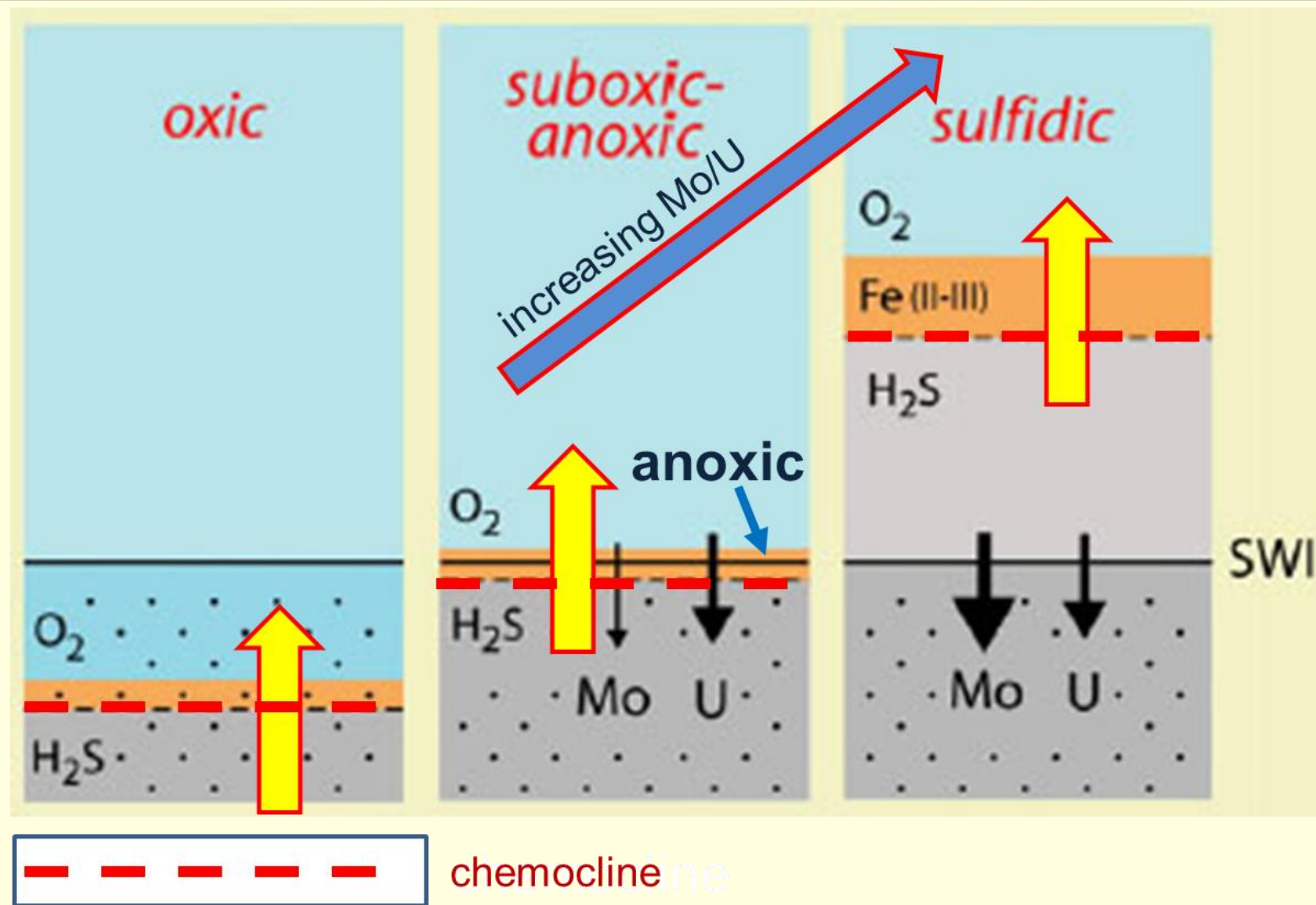




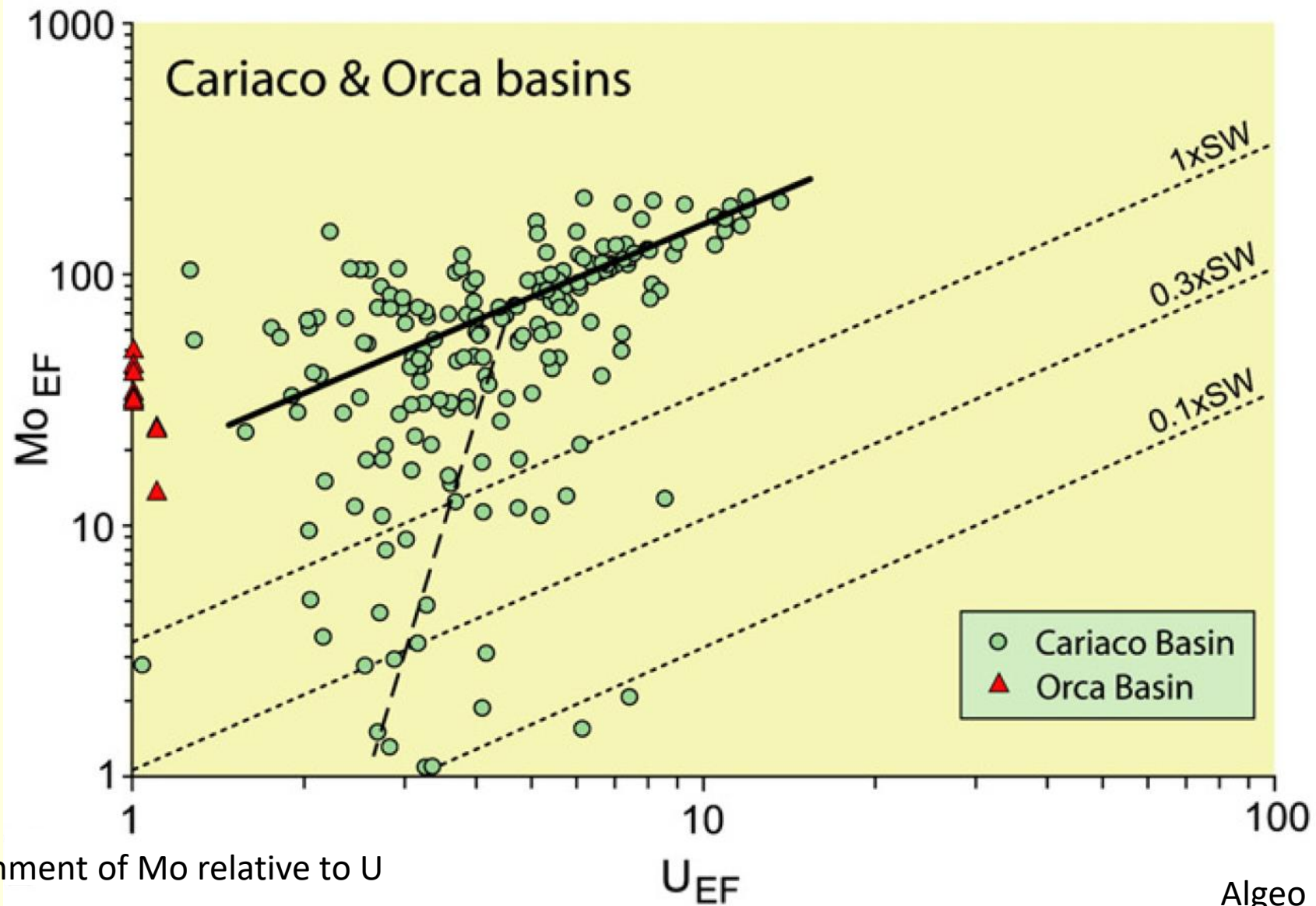
Mo depleted water mass (basin reservoir effect) → restricted basin



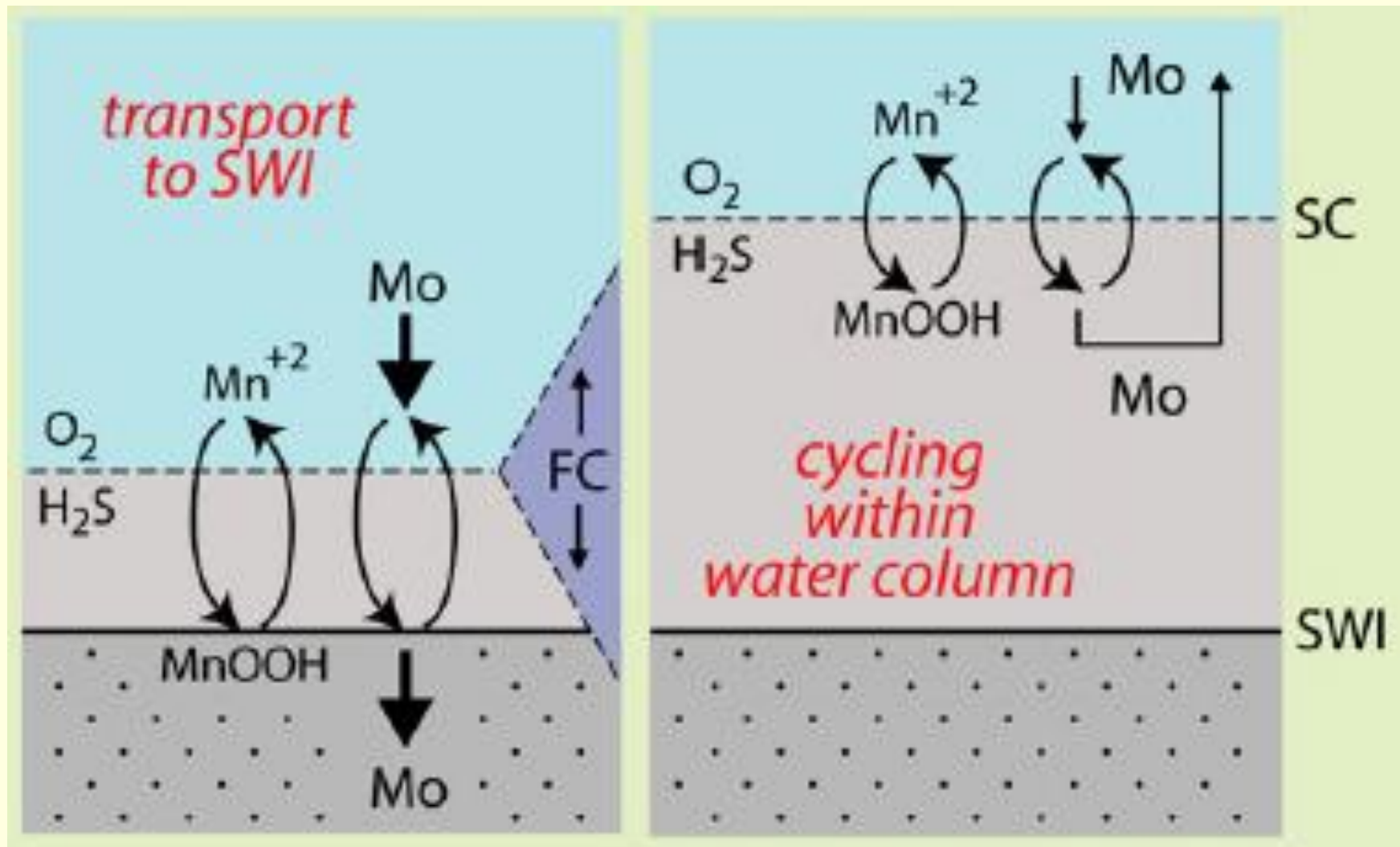
Progressive enrichment of Mo relative to U as total authogenic enrichment increases



Weakly restricted to open marine setting

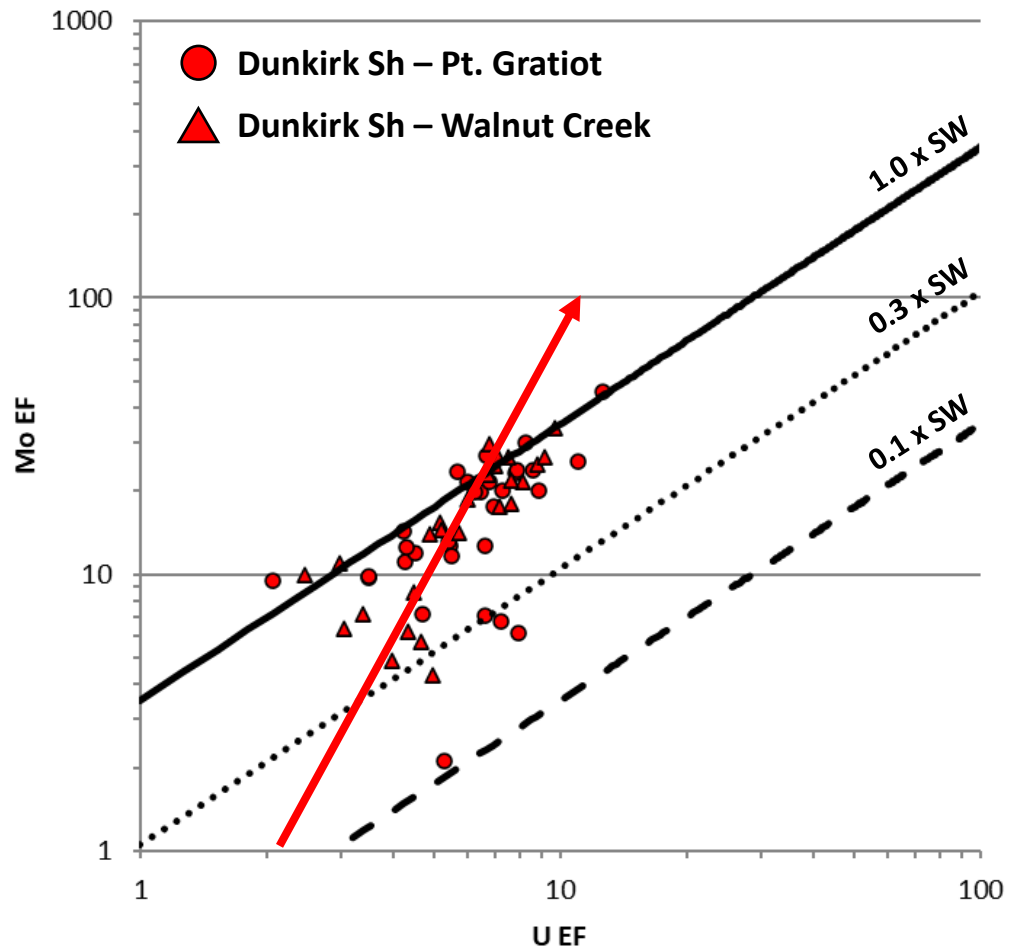


Considerable enrichment of Mo relative to U



Weakly restricted basin – requires intermittently sulfidic (dissolved H_2S) bottom water

Algeo and Tribovillard, 2009

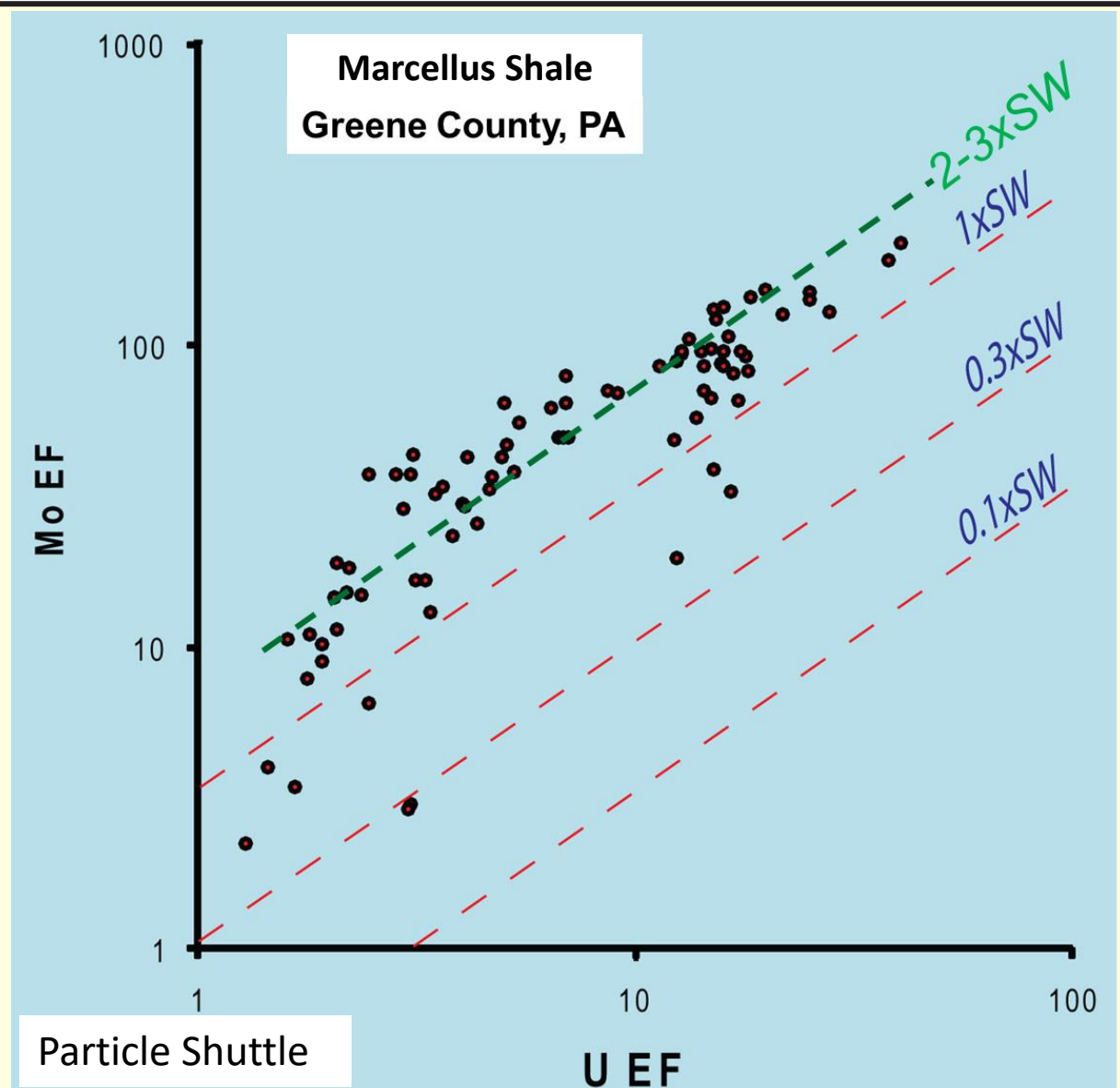
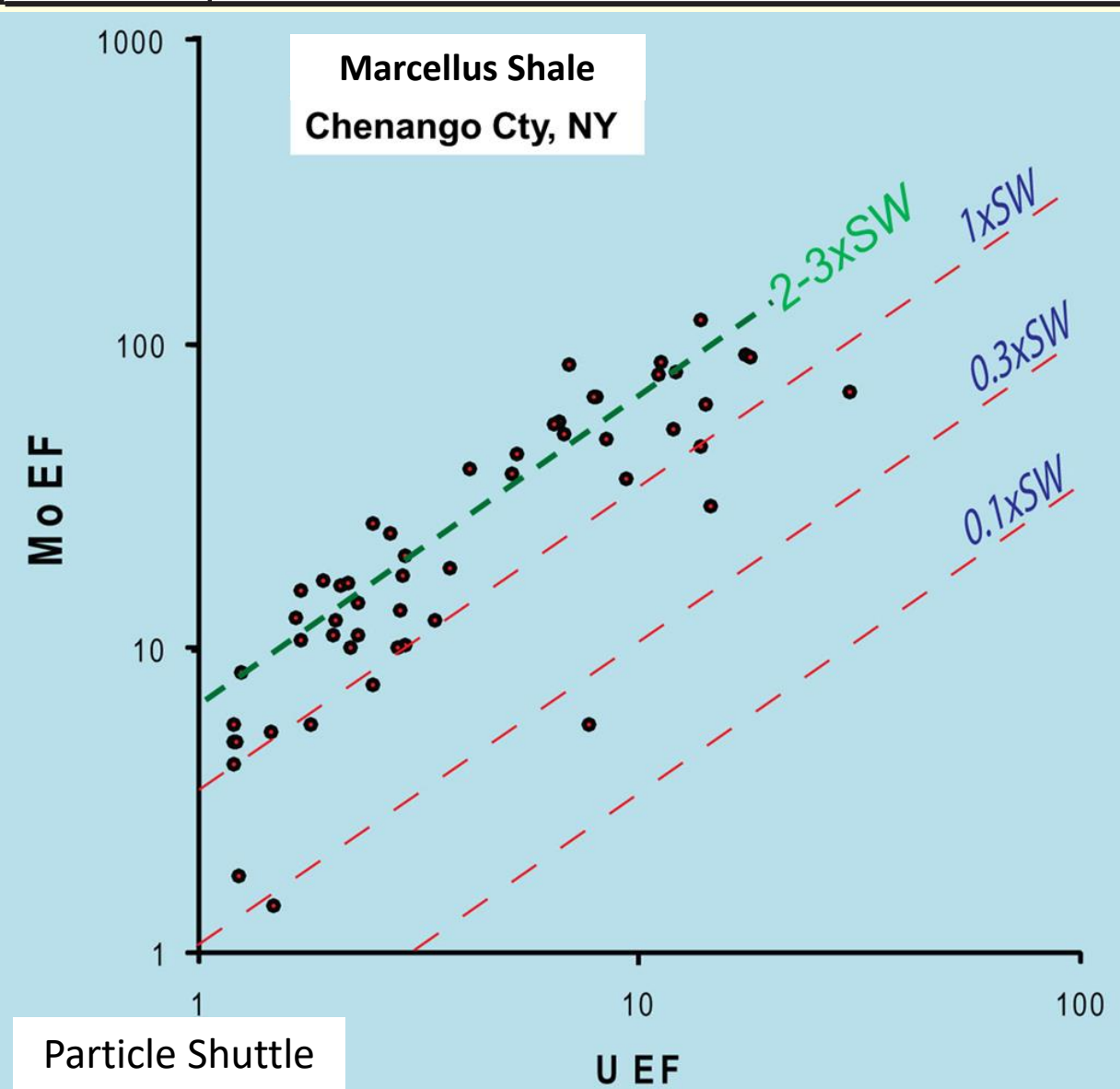


Weakly restricted basin

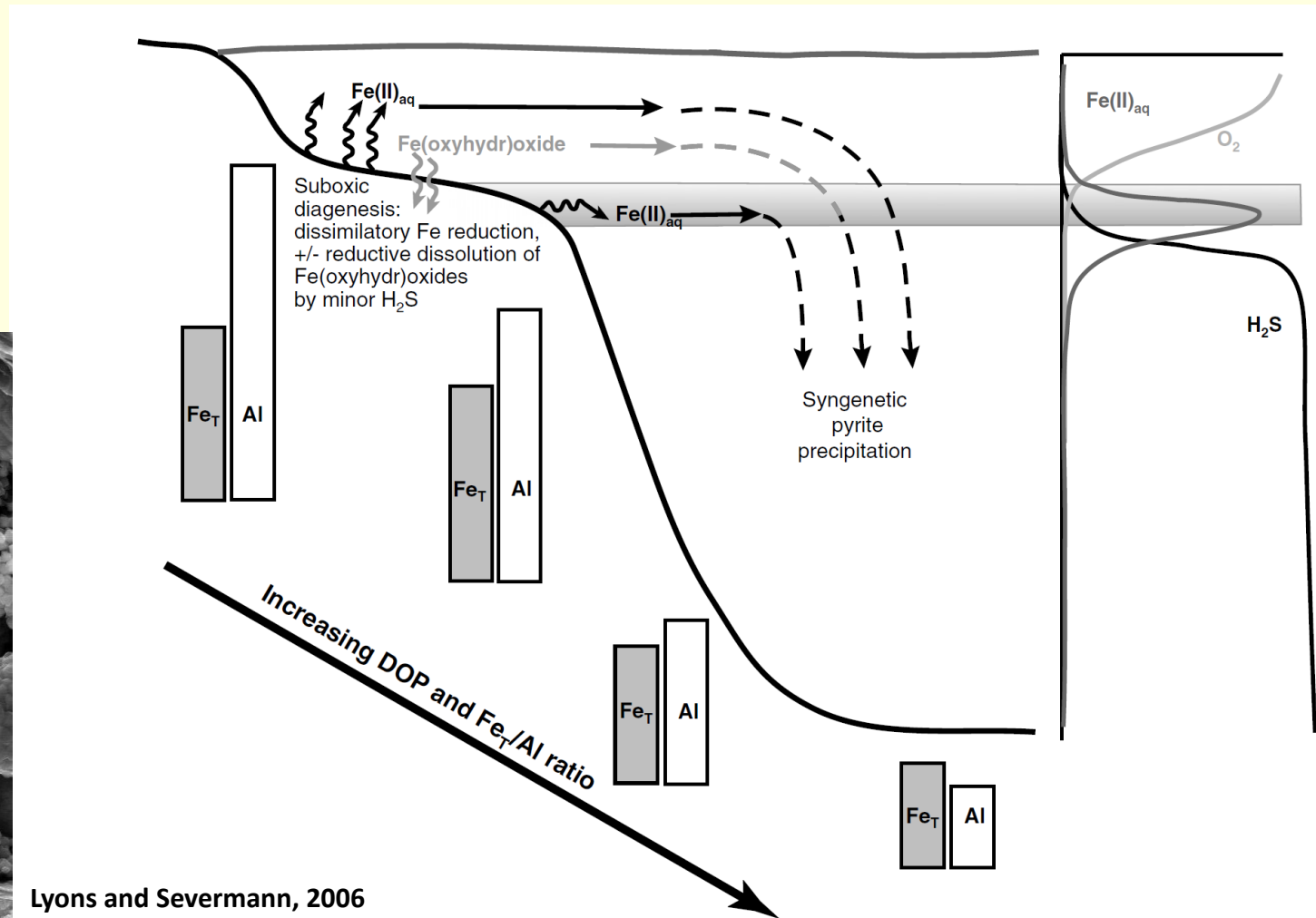
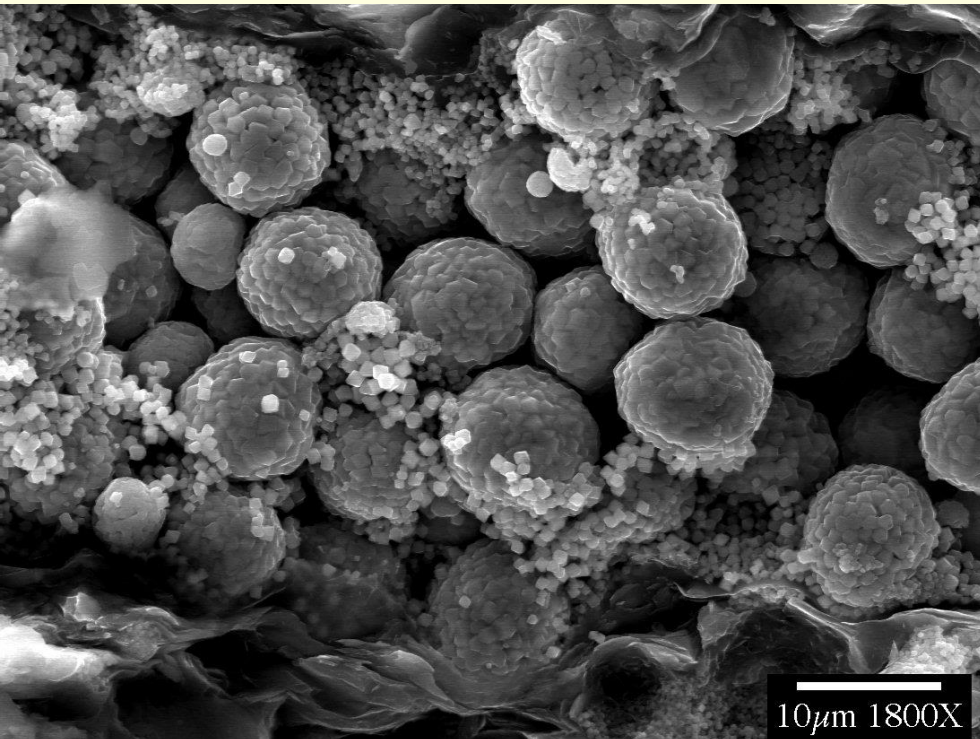


Uranium and Molybdenum profiles

Marcellus Shale



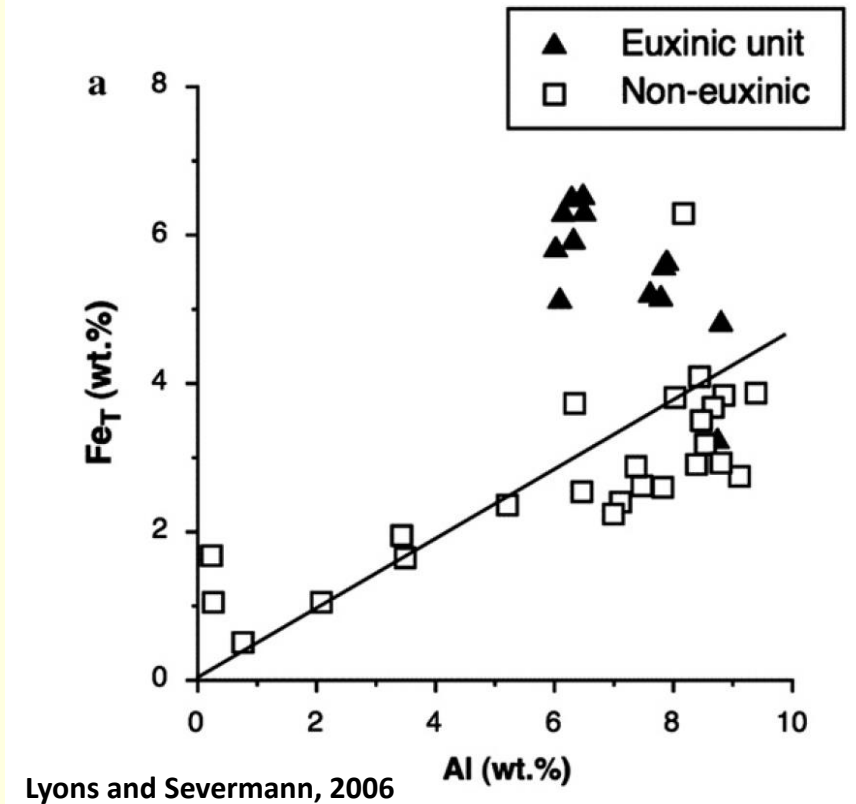
- The average shale value for Fe/Al is 0.55
- Diagenesis in shelf environments produce dissolved Fe II, some portion of which is exported to the bottom water
- In the euxinic bottom water, this Fe is sequestered into the sediment in the form of syngenetic/diagenetic pyrite



Lyons and Severmann, 2006

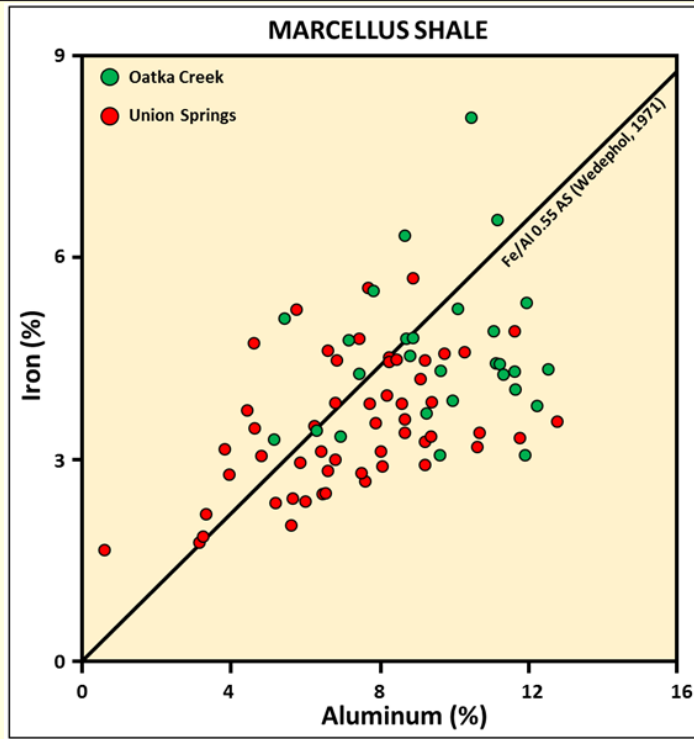
*DOP can be calculated from elemental data, but requires assumption that all S is in pyrite

- Two observations can be made about Fe/Al in euxinic settings:
 - Bacterial sulfate production is ubiquitous, therefore sedimentary pyrite concentrations and organic carbon are not linked to one another (although both are often elevated)
 - Under oxic water conditions sulfide production is limited to anoxic pore waters and therefore entirely dependent on the local availability of organic carbon.
- Fe/Al increase and Fe is no longer only a function of Fe delivery to the basin.
 - Fe/Al exceeds 0.55 and the trend is decoupled from Al.

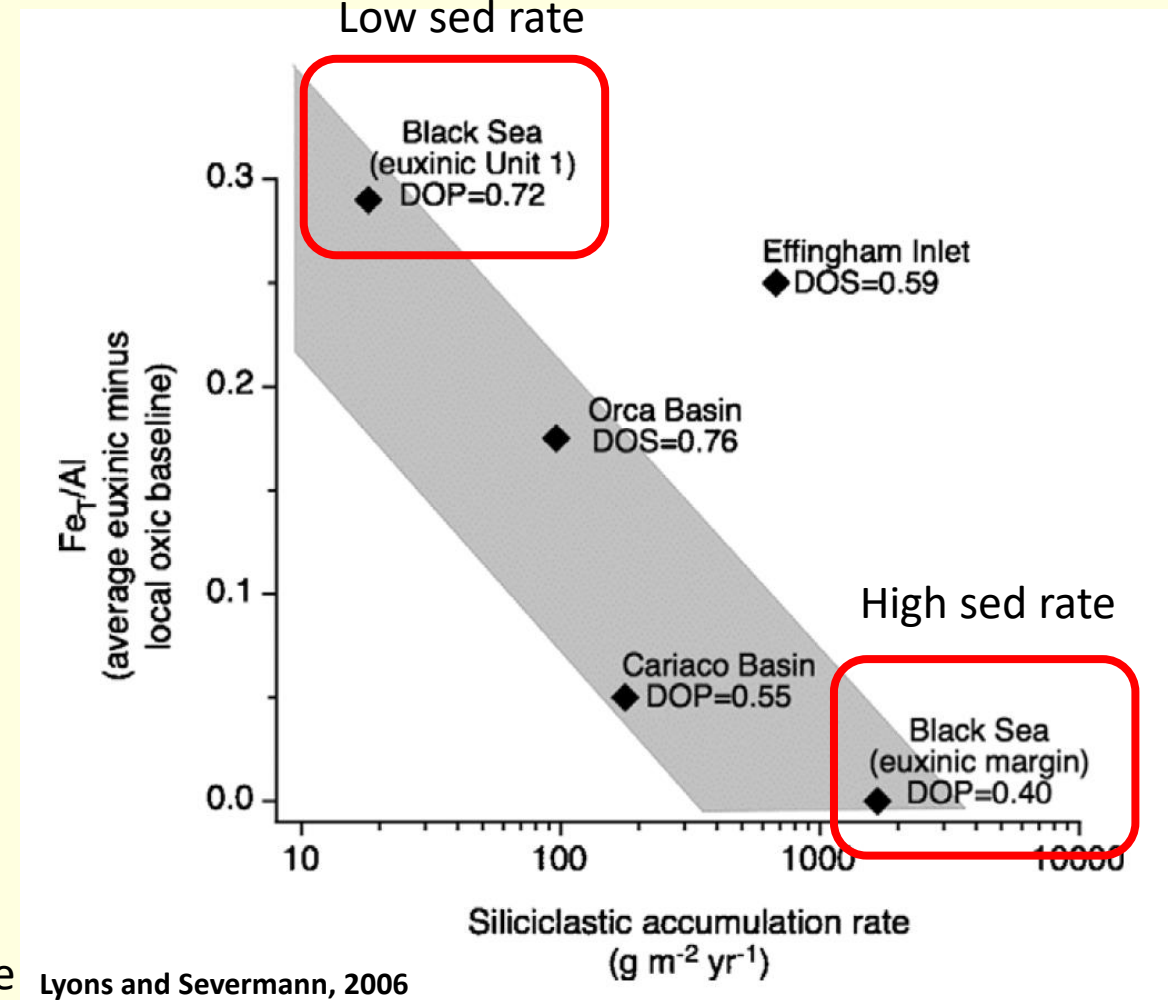


Interpretations that are not straight forward

Iron sequestration

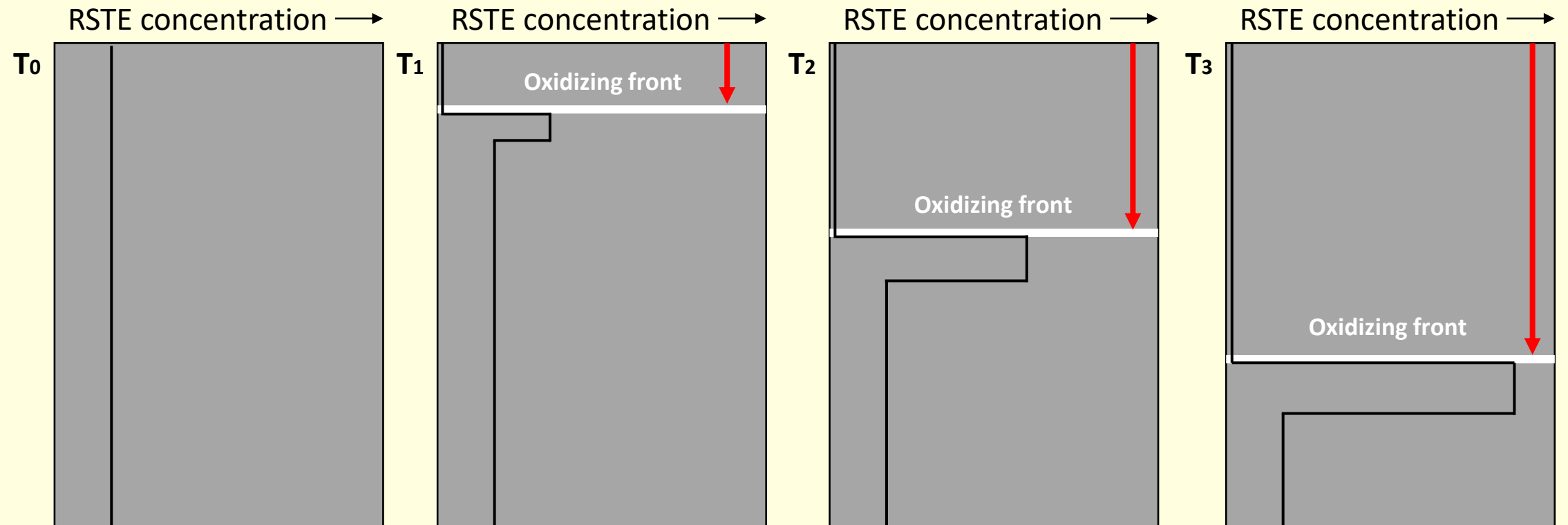


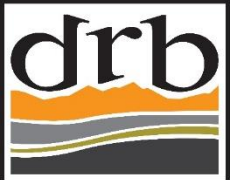
- Lack of correlation with Al argues against an oxic environment



- Must also consider sedimentation rate
- At high sedimentation rates, delivery of non-reactive Fe to the basin can exceed pyrite production leading to lower Fe/Al ratio
- The Marcellus accumulated over 1.0-1.5 my
 - Compare to the Barnett Shale → similar thickness yet accumulated over ~22 my...much lower sedimentation rate

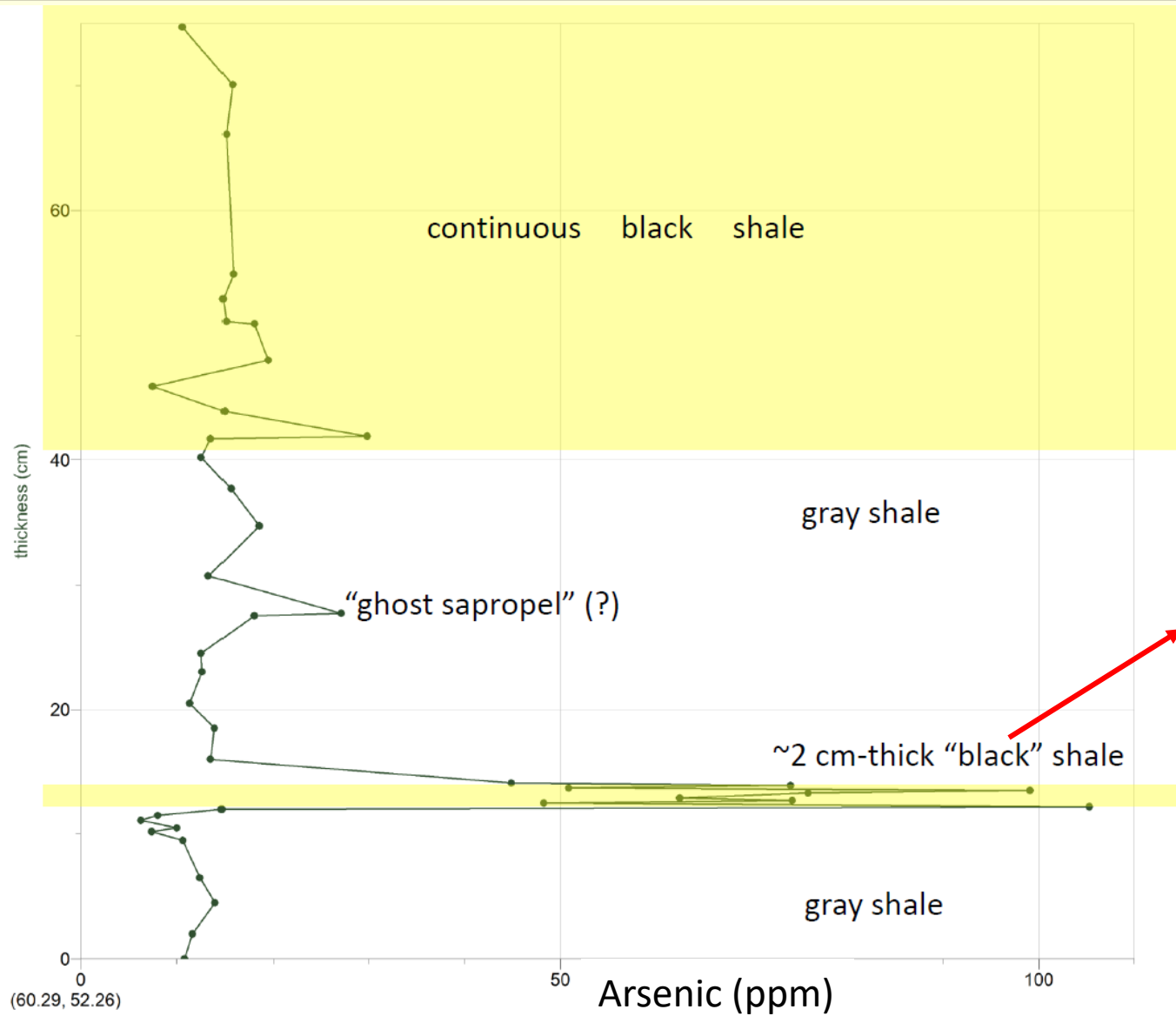
- Burn down has been offered as a way of explaining high concentrations of RSTE in black shales assumed to have been deposited under oxic conditions
- Burn down is a process by which an oxidizing front moves down through the sediment pile
 - This acts to dissolve and mobilize elements that are conservative under oxic conditions.
 - When the front stops moving these elements precipitate out in anoxic pores of sediment at higher concentrations





Remobilization of elements

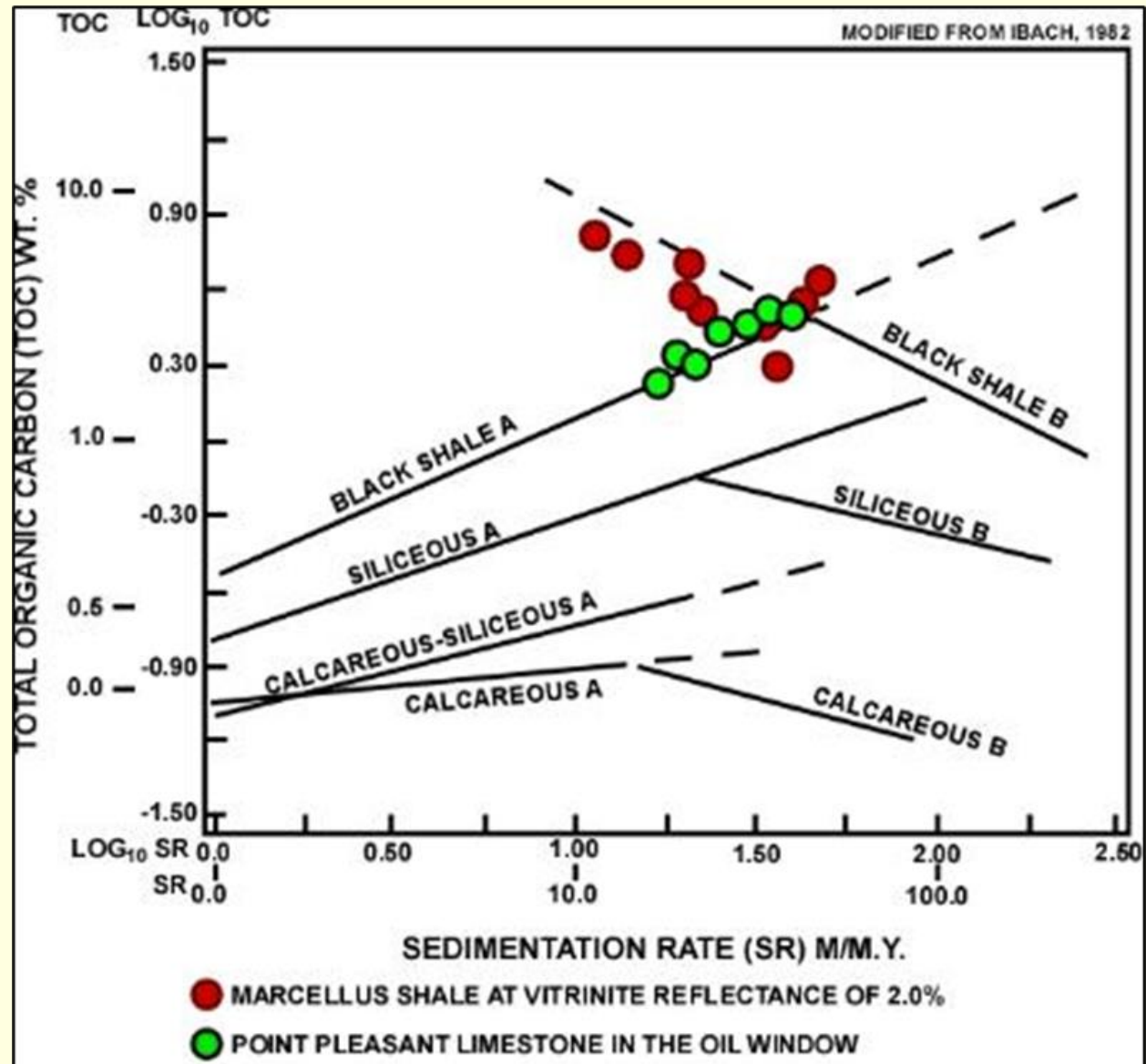
Burn Down



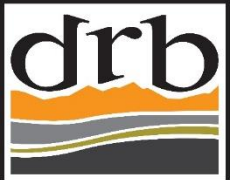


Purpose of the Work

Where is the reservoir?

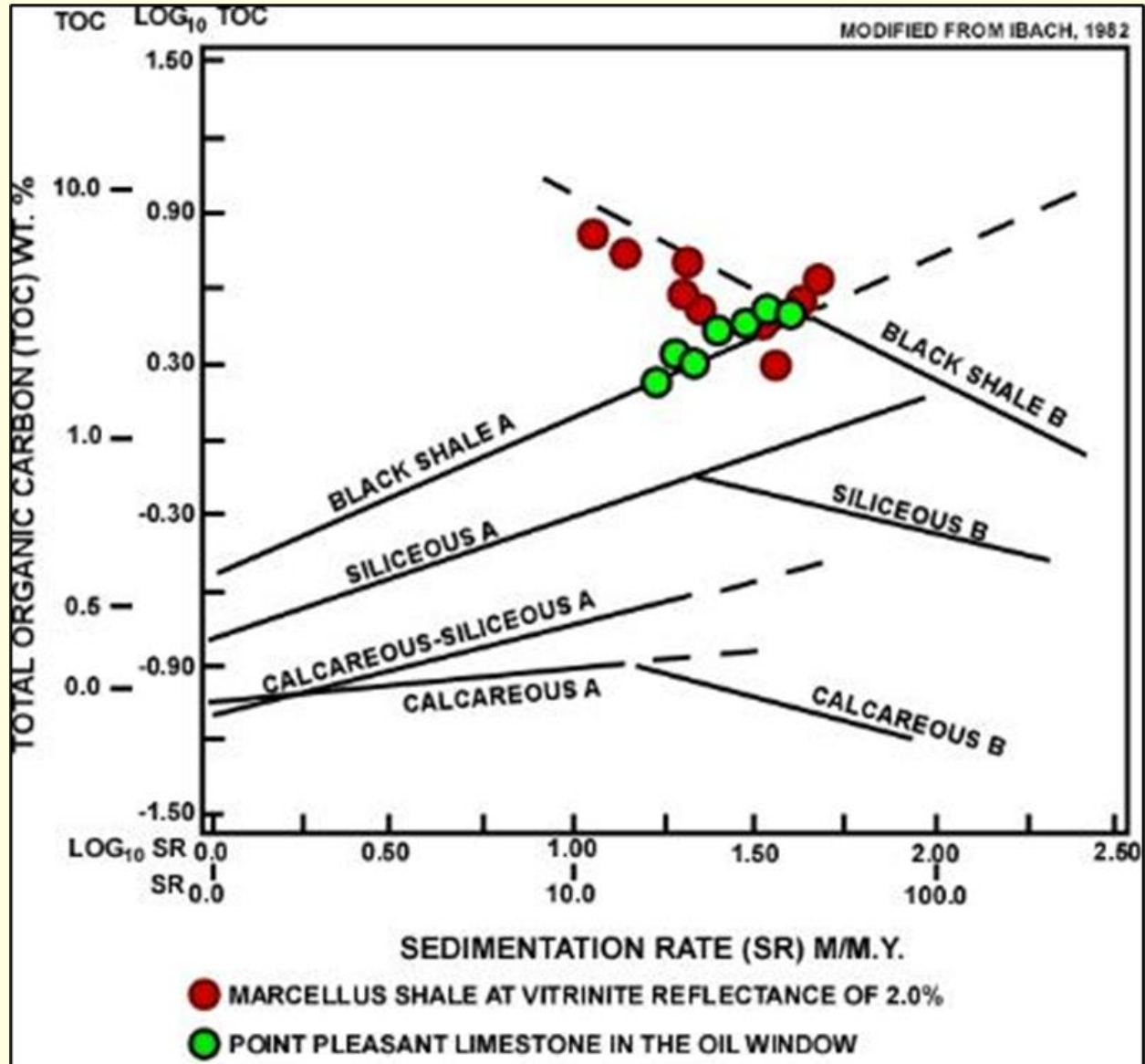


- Redox conditions can influence exploration, drilling, and completion techniques.
- Under oxic conditions, organic matter is preserved by rapid burial.
 - This removes organic material from zones of oxidation and scavenging by benthos
 - Under such conditions organic material is usually greatest in the thickest parts of the sedimentary pile, which may also include the greatest clay content.

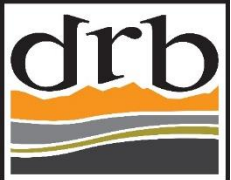


Purpose of the Work

Where is the reservoir?

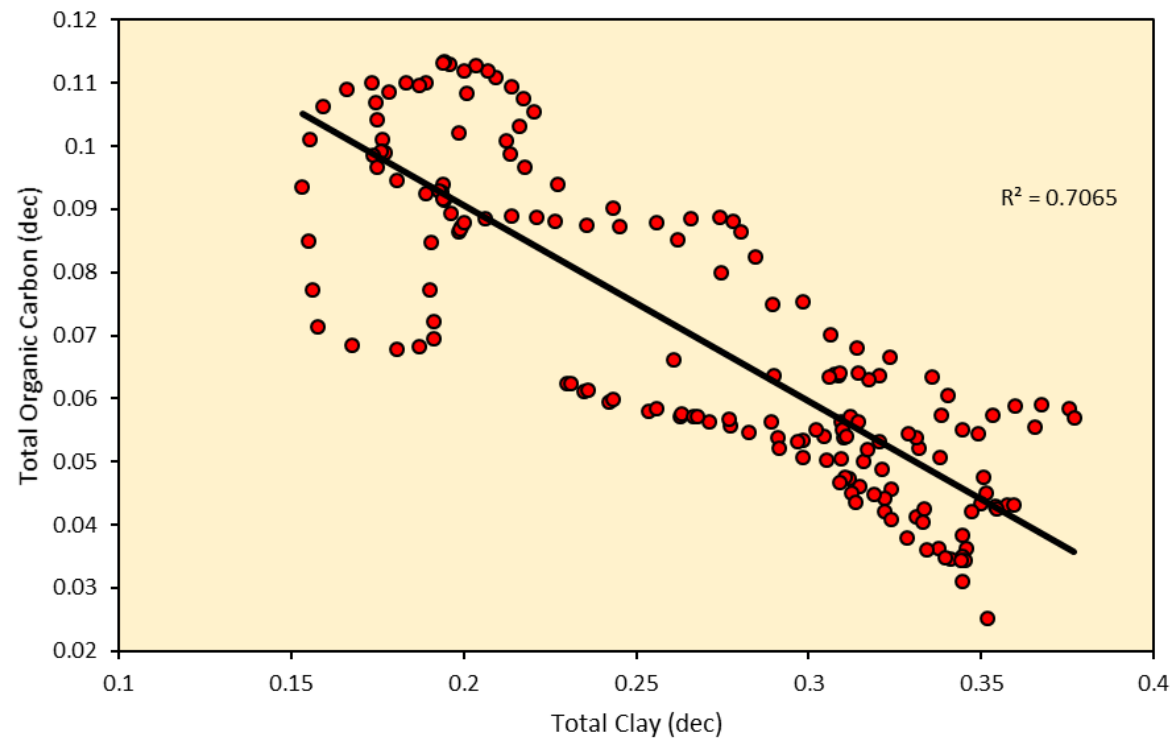


- Redox conditions can influence exploration, drilling, and completion techniques.
- Under anoxic conditions, dilution controls organic matter concentration.
 - A lack of oxygen and benthos increases the preservation potential of organic matter
 - If productivity is viewed as a constant, then organic matter concentration decreases with increasing sedimentation rate.



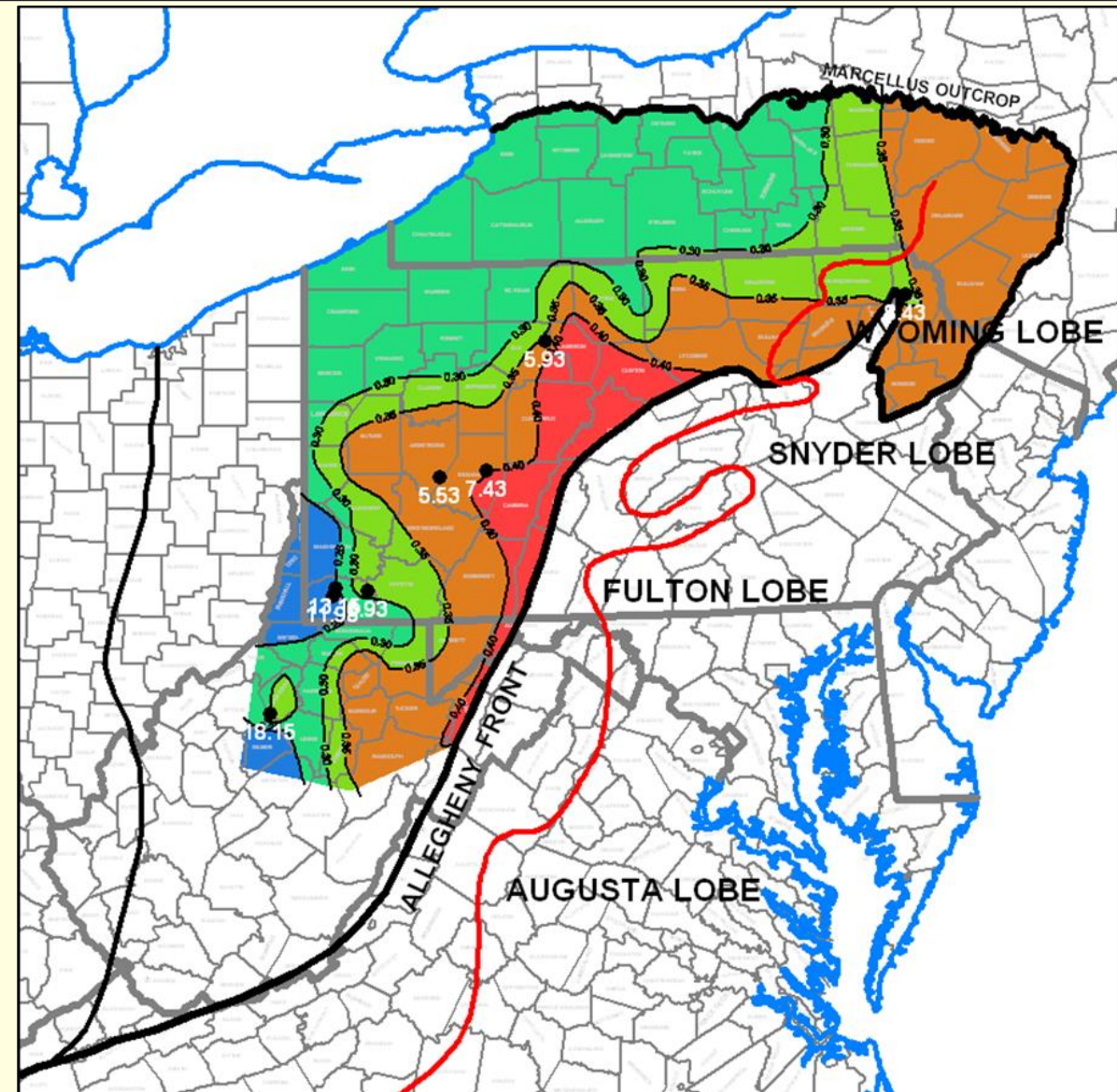
Depositional environment

Marcellus Shale



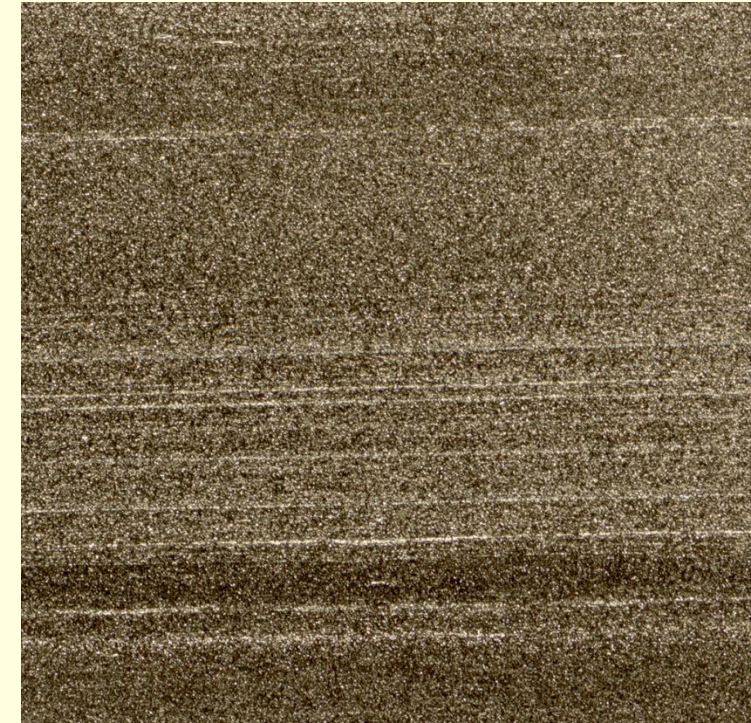
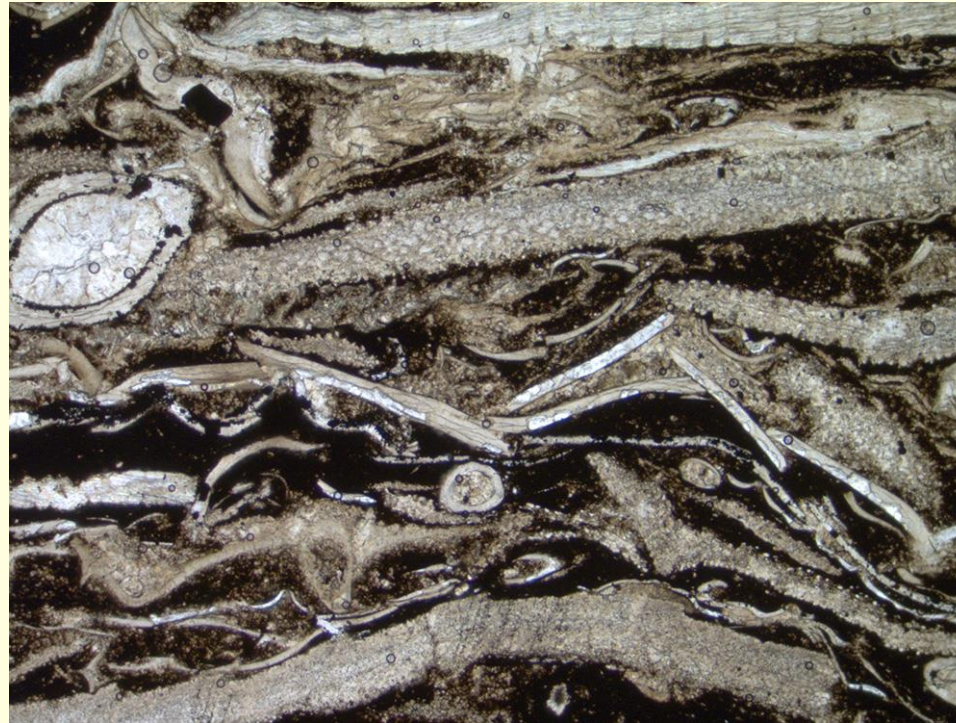
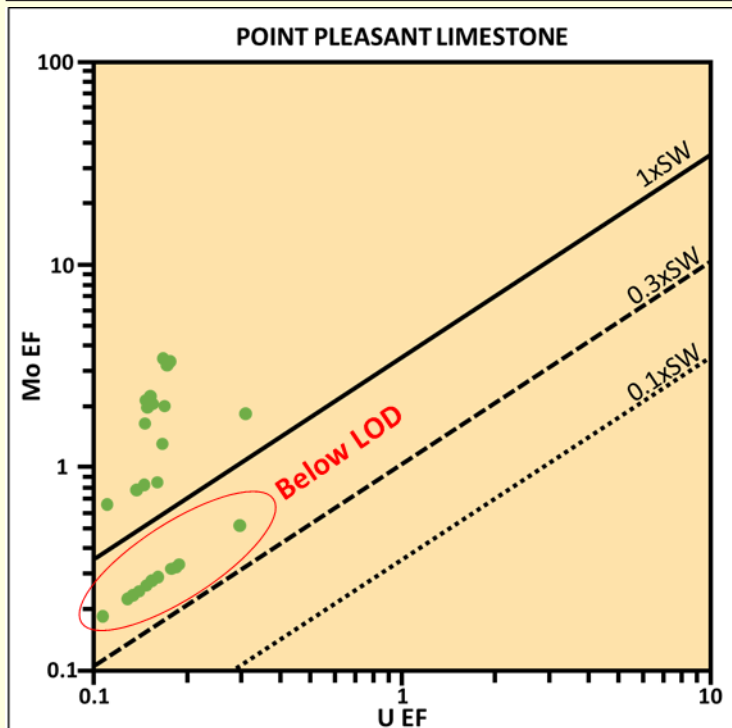
- TOC is inversely related to clastic influx

Organic matter concentration is controlled by dilution



(Dys)oxic organic-rich deposits

Point Pleasant Limestone

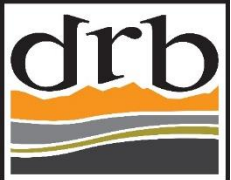


- Most RSTE below levels of detection
- Note scale change from previous plots (down one order of magnitude)

- Shell beds not transported
- Poorly sorted
- Articulated fossils present
- Grew *in situ*
- **Time-rich intervals**

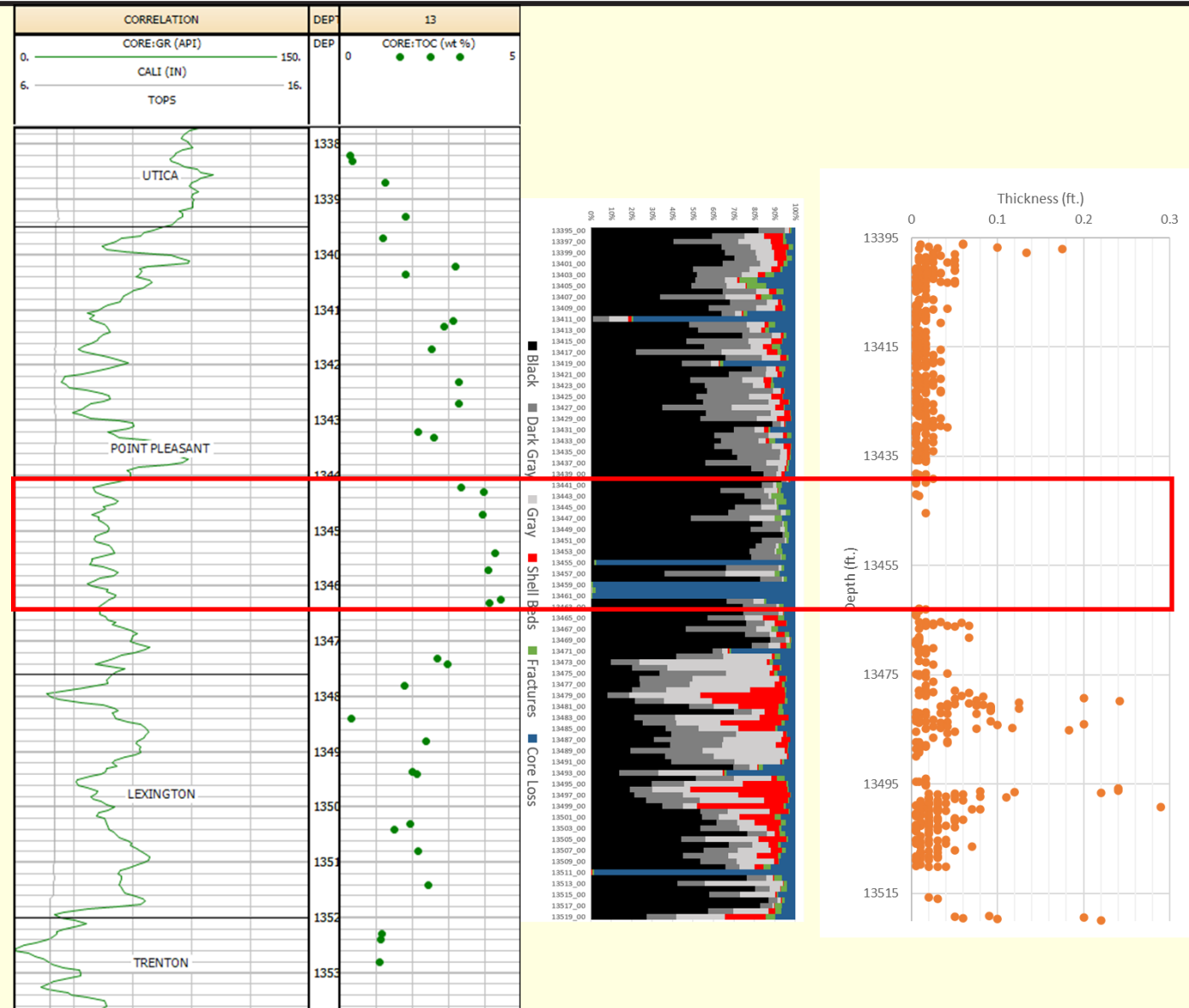
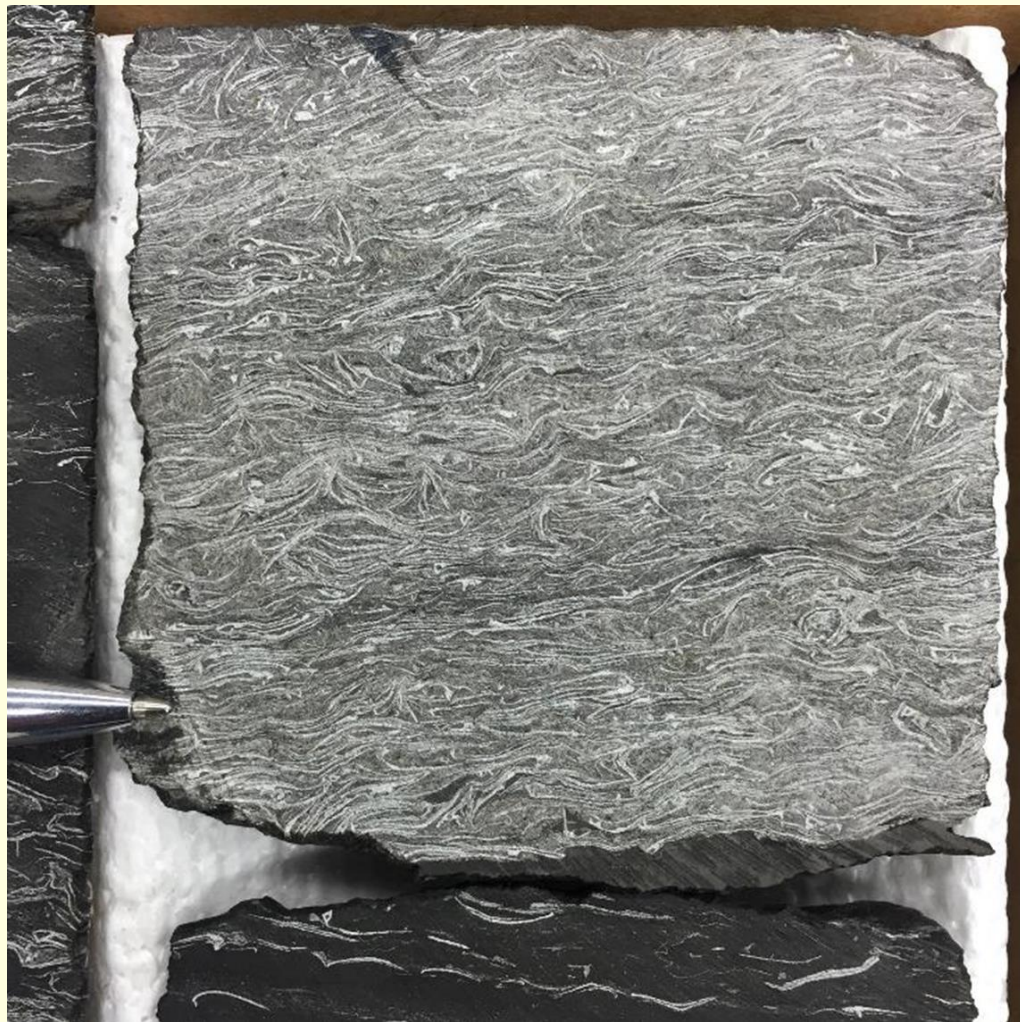
- Muds are well sorted
- Often show fining-up sequences
- Represent event beds
- Oubrution (suffocation) layers
- **Time-poor intervals**

Rapid deposition in an (dys)oxic environment led to the preservation of TOC



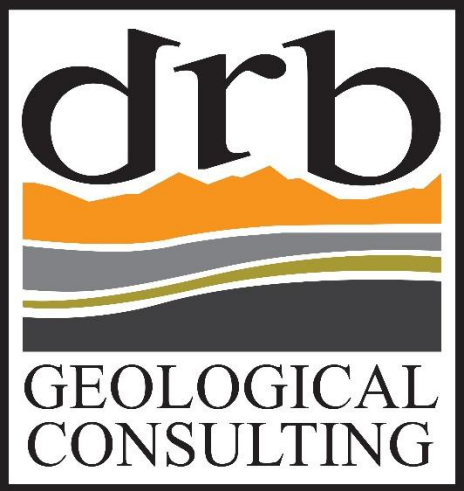
(Dys)oxic organic-rich deposits

Point Pleasant Limestone



- Not all interpretations are straight forward – integrate both elemental and sedimentological data sets
- Remember sample size – sidewall cores and core plugs can represent 100s to 1000s to 10,000s to even 100,000s of years
 - This results in a time-averaged interpretation of the rock record
- Elemental relationships are not universal. For example, a correlation between uranium and TOC in one shale may not be applicable or exist in another shale.





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